

Diet and habitat of unique individuals of *Dinohippus mexicanus* and *Neohipparrison eurystyle* (Equidae) from the late Hemphillian (Hh3) of Guanajuato and Jalisco, central Mexico: stable isotope studies

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ABSTRACT

Stable carbon and oxygen isotopes were determined in molar enamel from fossil Pliocene equids from Rancho El Ocote in the San Miguel Allende basin, Guanajuato, and from Santa María, Tecolotlán basin, Jalisco. At each locality, the source was one molar from an individual *Dinohippus mexicanus* and one molar from an individual *Neohipparrison eurystyle*. Results indicated that the *N. eurystyle* individuals from both localities had been C₃/C₄ mixed feeders, and had lived in open-zone vegetation ($\delta^{13}\text{C}$: -3.1‰ to -1.3‰; $\delta^{18}\text{O}$: -4.9‰ to -6.4‰). On the other hand, the *D. mexicanus* from Rancho El Ocote had fed upon C₄ plants and lived in open zones ($\delta^{13}\text{C}$: -1.3‰; $\delta^{18}\text{O}$: -4.9‰), whereas the *D. mexicanus* from Santa María was a C₃/C₄ mixed feeder with considerable consumption of C₃ plants ($\delta^{13}\text{C}$: -7.7‰; $\delta^{18}\text{O}$: -6.4‰). These results could be contrasted to suggestions from previous isotopic work that *D. mexicanus* in Mexico predominantly fed on C₄ plants and further samples analyses are warranted. This study contributes to the understanding of the Pliocene equid taxa from central Mexico and emphasizes the presence of different diets, ranging from exclusive C₄ to mixed C₃/C₄ plants.

Key words: *Dinohippus mexicanus*; *Neohipparrison eurystyle*; stable isotopes; late Hemphillian, Mexico.

RESUMEN

En este trabajo se analizó un molar de *Dinohippus mexicanus* y uno más de *Neohipparrison eurystyle*, ambos recolectados de la localidad de Rancho el Ocote, en la Cuenca de San Miguel de Allende; de la misma forma también se analizaron un molar de cada una de las especies de équidos antes mencionadas, que fueron recolectados en la localidad de Santa María, de la cuenca de Tecolotlán, Jalisco. Los molares fueron analizados por isótopos estables de carbono y oxígeno presentes en el

esmalte dental, los cuales revelan que en ambas localidades, *N. eurystyle* poseía una dieta mixta C₃/C₄ y habitaba en zonas de vegetación abierta ($\delta^{13}\text{C}$: -3.1‰ y -1.3‰; $\delta^{18}\text{O}$: -4.9‰ y -6.4‰). En cambio, *D. mexicanus* de Rancho El Ocote se alimentó de plantas C₄ y vivió en zonas abiertas ($\delta^{13}\text{C}$: -1.3‰; $\delta^{18}\text{O}$: -4.9‰), mientras que el ejemplar de Santa María poseía una dieta mixta C₃/C₄ con un importante consumo de plantas C₃ ($\delta^{13}\text{C}$: -7.7‰; $\delta^{18}\text{O}$: -6.4‰), lo que es contrario a otras propuestas basadas en isótopos estables que indican que los *D. mexicanus* de México sólo se alimentaban de plantas C₄, lo cual contribuye a comprender que los hábitos alimentarios de las especies de équidos que habitaron durante el Plioceno el centro de México variaba de animales con consumo exclusivo de plantas C₄ a aquellos con dietas mixtas C₃/C₄.

Palabras clave: *Dinohippus mexicanus*; *Neohipparrison eurystyle*; Hemifiliano tardío; isótopos estables; México.

INTRODUCTION

Taxonomic and biostatigraphic research in the Mexican Late Tertiary sedimentary basins has shown a highly diversified mammal fauna. Among the best known and important local faunas are those from Rancho El Ocote in the San Miguel Allende basin, Guanajuato, and from Santa María in the Tecolotlán basin, Jalisco. The faunal complex known for those late Hemphillian localities includes the following: equids, *Nannippus aztecus*, *Neohipparrison eurystyle*, *Astrohippus stockii* and *Dinohippus mexicanus*; rhinoceros, *Teleoceras fossiger*; carnivores, *Agriotherium schneideri*, *Borophagus secundus*, *Canis ferox*, and *Machaerodus cf. M. coloradensis*; camelids, *Hemiauchenia vera*, *Megatylopus matthewi*, and *Alforjas* sp.; rodents, *Spermophilus* sp., *Ammospermophilus* sp., *Paenemarmota* sp., *Perognathus* sp., *?Pliogeomys* sp., *Calomys* sp., *Baiomys* sp., *Prosigmodon* sp. and *Neotoma* sp.; the xenarthran *Megalonyx* sp.; lagomorphs, *Hypolagus mexicanus* and *Notolagus velox*; and the proboscidean *Rhynchotherium* sp. (Miller

and Carranza-Castañeda, 1984, 1998; Carranza-Castañeda, 2006; Jiménez-Hidalgo and Carranza-Castañeda, 2010; Carranza-Castañeda *et al.*, 2013). These faunas constitute the basis for the late-Hemphillian biostratigraphy of central Mexico.

Carbon isotope values from dental enamel record the photosynthetic pathways of plants consumed by the animal in an early life stage when the molars were being formed (Koch, 2007). On the other hand, the mesowear method is based on the attrition (tooth-on-tooth contact) and abrasion (tooth-on-food contact) of the teeth (largely their cusps), regardless of the photosynthetic pathway used by the plant that generates it (Kaiser and Solounias, 2003). This method assesses the dental wear caused by the animal's diet during months to years of its life span, *i.e.*, the long-term alteration of the enamel (Fortelius and Solounias, 2000). The abrasion produced by the intake of abiotic material (*e.g.* sand, dust, or soil) could be many times more severe than the effect of enamel attrition (Sansou *et al.*, 2007).

Analysis of carbon and oxygen stable isotopes in fossil molars of the equids *Astrohippus stocki*, *Dinohippus mexicanus*, *Nannipus aztecus*, and *Neohipparrison eurystyle* from Yépómera, Chihuahua, and Rancho El Ocote, Guanajuato, suggested that these species fed predominantly on C₄ plants, such as grasses or other herbaceous plants (MacFadden, 2008). On the other hand, a mesowear analysis of molars from *D. mexicanus* and *N. eurystyle* from the Tecolotlán basin (Jalisco) indicated that *N. eurystyle* had specialized in feeding on abrasive material (such as grass or external grit), whereas *D. mexicanus* was a mixed feeder (Barrón-Ortiz and Guzmán, 2008). A combination of mesowear, micro-wear and stable isotope analyses (Bravo-Cuevas *et al.*, 2015) suggested that *A. stocki* from Rancho El Ocote had a C₃/C₄ mixed diet; most C₃ plants are shrubs and trees, with a few C₄ grasses. Unfortunately, those authors did not provide the stratigraphic context of their specimens. Carbon and oxygen stable isotopes in molars of *D. mexicanus* and *N. eurystyle* from the Tecolotlán basin have not previously been analyzed. In the present study, enamel from the molar of an individual from each species from Tecolotlán basin was analyzed and the data were compared with those from the molar of an individual of each species from Rancho El Ocote. The objective was to detect differences and similarities in diet of the two species in comparison with the results recorded by MacFadden (2008).

δ¹³C and δ¹⁸O biogeochemical markers

Since 1980, the application of carbon and oxygen stable isotopes analysis has been important in inferring diet and habitat of Cenozoic herbivorous mammals and birds (Berrick, 1998; Koch, 1998; MacFadden and Higgins, 2004; Angst *et al.*, 2014). Carbon is incorporated into animals by three plant photosynthetic pathways. The C₃ pathway has δ¹³C from -22‰ to -35‰ and is found in trees and shrubs, as well as in some cold-zone grasses (Smith and Epstein, 1971; O'Leary, 1981; Medina *et al.*, 1986). The C₄ pathway has δ¹³C values from -10‰ to -12‰ and is typical for grasses, and also for some dry-zone trees and shrubs (Stowe and Teeri, 1978; Cerling, 1999; Cerling and Ehleringer, 2000). Finally, the Crassulacean Acid Metabolism (CAM) pathway has δ¹³C values from -35‰ to -12‰ and is found in succulent plants such as cacti and some orchids (Gröcker, 1997; Andrade *et al.*, 2007).

Herbivores incorporate the carbon of the plants on which they feed into their organs and tissues, so they have carbon isotope values reflecting those of the plants they feed upon. In the case of tooth enamel, the tissue is enriched by 14.1‰ relative to the plant value (Cerling and Harris, 1999). Animals that feed on C₃ plants have carbon isotopic values from -19‰ to -9‰, while herbivores that have consumed C₄ plants have carbon isotopic values from -2‰ to +2‰. The C₃/C₄ mixed feeders show values from -9‰ to -2‰ (MacFadden and Cerling, 1996).

On the other hand, oxygen is incorporated into animals by inhala-

tion, from water in food, and mainly by ingested water. Such oxygen is in equilibrium with what is lost through CO₂ exhalation, feces, urine and sweat (Koch *et al.*, 1994; Sánchez, 2005). Given that ingested water comes from meteoric water, its composition is affected primarily by temperature, but also by altitude, latitude and the amount of rain fall in a zone (Dansgaard, 1964). The oxygen composition of dental enamel from mammals is used mainly for inferring the paleoclimatic conditions that prevailed in a locality in the past (Longinelli and Nuti, 1973; Kohn, 1996; Kohn *et al.*, 1998; Schoeninger *et al.*, 2000; Mahboubi *et al.*, 2014).

STUDY LOCALITIES

Rancho El Ocote is located at 21°05'28"N and 100°41'01"W in the San Miguel Allende basin, Guanajuato. The ashes found in the stratigraphic sequence were dated by fission tracks, and provided ages of 4.8 Ma for the lower sediments (the Rhino layer) formed during the late Hemphillian, and 4.7 Ma for the upper ash from the early Blancan (Kowallis *et al.*, 1998). These ages provide the limits of the North American Land Mammal Ages (NALMA), late Hemphillian and early Blancan (Figures 1 and 2) (Flynn *et al.*, 2005; Carranza-Castañeda, 2006; Cohen *et al.*, 2013). Four equid species, *D. mexicanus*, *N. eurystyle*, *A. stockii*, and *N. aztecus* occur in the Rhino level, with lithology and sedimentation rates showing a more humid environment than those of the overlying Blanco level, where only *D. mexicanus* and *N. neurystyle* are found (Carranza-Castañeda, 2006).

The molars from Rancho El Ocote were collected at the GTO-2A locality. Carranza-Castañeda *et al.* (2013) recorded that *Dinohippus mexicanus* and *Neohipparrison eurystyle* were found in the Rhino level, and to Hh3 in the Mammalian Chronology proposed by Tedford *et al.* (2004). On the other hand, *N. eurystyle* was not found above the Rhino level. This implies that the level above the Rhino level represents the final portion of the late Hemphillian (Hh4) (Figure 2).

The Tecolotlán basin is 100 km southeast of Guadalajara, in the state of Jalisco (Figure 1; the late Hemphillian stratigraphic sequence is divided by lacustrine sands). The best faunal representation is from the Santa María locality and the studied molars were collected at this site (Figure 3). Lower ash has been dated by ⁴⁰Ar/³⁹Ar at 4.89 Ma (Kowallis *et al.*, 1998). Carranza-Castañeda *et al.* (2013) have shown that the Santa María area belongs to Hh3. As an erosional discordance, the late Blancan-Irvingtonian, Buenaventura sequence contains gravel and sands, with an internal ash dated at 2.6 Ma (Kowallis *et al.*, 1998). Santa María has a stratigraphic sequence in the Tecolotlán basin that terminates at the bottom of the lacustrine sands. Lithology and sedimentation rates at Santa María indicate the presence of a humid environment, containing remains of *D. mexicanus*, *N. eurystyle*, *A. stockii* and *N. aztecus*. The last two species are not found in the San José stratigraphic sequence that is overlying the lacustrine sands, with lithology and sedimentation rates typical for a dry environment (Carranza-Castañeda *et al.*, 2013).

MATERIALS AND METHODS

Sample extraction, preparation and statistical analysis of the results

A bulk sample was taken from isolated teeth from Rancho El Ocote: MPGJ-2056 (*Dinohippus mexicanus*), MPGJ-2026 (*Neohipparrison eurystyle*) and Tecolotlán Basin: MPGJ-2956 (*D. mexicanus*), MPGJ-2959 (*N. eurystyle*), each of them representing a single individual. The samples were taken at the crown-root joint following the protocol of



Figure 1. Late Hemphillian fossiliferous localities in the Tecolotlán basin (1) and Rancho El Ocote (2), Mexico.

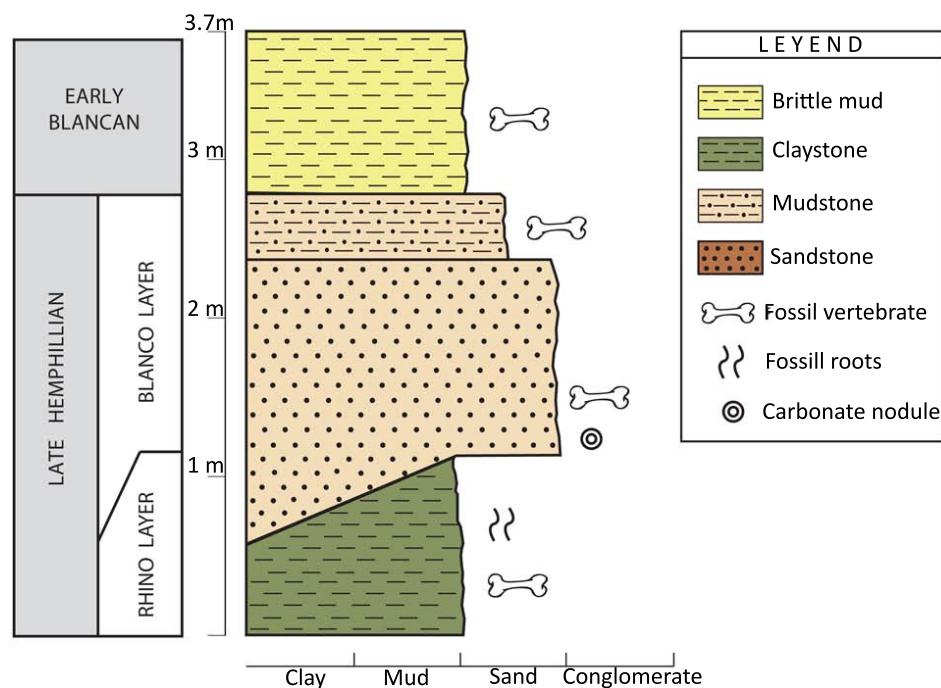


Figure 2. Stratigraphic column for Rancho El Ocote (GTO 2A). Investigated specimens came from the Rhino layer. Taken from Pacheco-Castro and Carranza-Castañeda (2015).

Feranec and MacFadden (2000) and belong to the Palaeontological Collection of the *Centro de Geociencias* of the *Universidad Nacional Autónoma de México* (UNAM), *Campus Juriquilla, Querétaro*.

Samples were processed in the Stable Isotope Laboratory at the *Instituto de Geología*, UNAM, with the method proposed by Koch *et al.* (1997). First, 20 mg of enamel was ground and sieved (125 μ m mesh) to obtain a fine and uniform powder. Then 10 ml of hydrogen peroxide at 30 % was added to eliminate the organic matter. After 2 h, the samples were centrifuged and the hydrogen peroxide decanted and washed again three times with water type I (grade HPLC 18.2 M Ω). Once the washing was finished, 5 ml of a buffer solution, $\text{Ca}(\text{CH}_3\text{CO}_2)_2\text{-CH}_3\text{COOH}$ 1.0 M, pH 4.75, was added and allowed to rest for 9 h. The buffer solution was decanted and the samples were washed again three times with water type I. Finally, to eliminate any remaining water, ethanol was added, and the solution was left for 20 h in an oven at 90 °C. Isotopic ratios were determined with a Finnigan MAT 253 mass spectrometer with a dual inlet system and auxiliary Gas Bench equipment with a GC Pal autosampler with a temperature-controlled aluminum plate adjoined to the mass spectrometer (Révész and Landwehr, 2002). Results were reported as $\delta^{18}\text{O}_{\text{VPDB}}$ and $\delta^{13}\text{C}_{\text{VPDB}}$, and they were normalized using NBS-19, NBS-18 and LSVEC to the Vienna Pee Dee Belemnite (VPDB) scale in accordance with the corrections described by Coplen (1988), Werner and Brand (2001) and Coplen *et al.* (2006). For this technique, the standard deviation was 0.2‰ for oxygen and carbon.

RESULTS

Carbon and oxygen isotopic values for *D. mexicanus* from Rancho El Ocote are more positive than those from Tecolotlán (Table 1, Figure 4). The ^{13}C values for specimens of *N. eurydice* from the two sites are similar to each other, although somewhat higher from Rancho El Ocote than from Tecolotlán (Table 1, Figure 4).

DISCUSSION

Diet

The carbon isotope value ($\delta^{13}\text{C}$: -1.3‰) for *Dinohippus mexicanus* (MPGJ-2056) from Rancho El Ocote implies that this individual was a specialist feeding on C_4 plants such as grasses or other herbs (Table 1), also observed by MacFadden (2008) for Rancho El Ocote and Yepómera locations. In contrast, individuals found in Florida and Texas were mixed C_3/C_4 feeders (MacFadden, 2008), but consumed an important amount of C_3 plants, such as C_3 grasses, tree leaves and shrubs. On this basis, MacFadden (2008) assumed that, during the late Hemphillian, C_4 plants were abundant in Mexican ecosystems, whereas C_3 plants were abundant further north. However, the present results with specimen MPGJ-2956 from Santa María (Tecolotlán basin) differ in that they show a $\delta^{13}\text{C}$ value that characterize it as a C_3/C_4 mixed feeder, but with an important component of C_3 plants (Table 1).

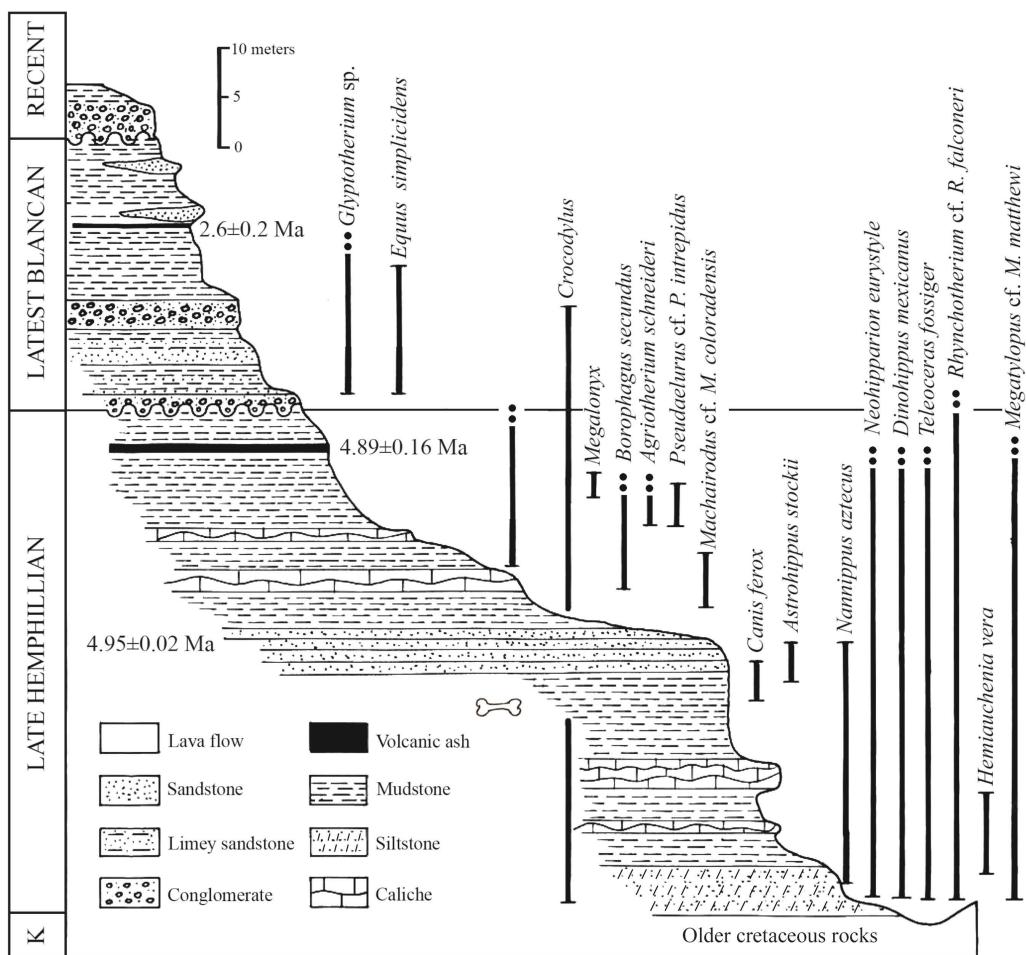


Figure 3. Stratigraphic column for the Tecolotlán basin (Santa María). Investigated specimens came from the bone layer. B: Bone layer, K: Cretaceous. Taken from Carranza-Castañeda (2006).

Table 1. Carbon and oxygen isotopic values of molar enamel from fossil equids in Mexico.

Catalogue number	Species	$\delta^{13}\text{C}_{\text{VPDB}\%}$	$\delta^{18}\text{O}_{\text{VPDB}\%}$	Locality	Diet suggested by MacFadden (2008)	Present study
MPGJ-2056	<i>Dinohippus mexicanus</i>	-1.3	-3.3	Rancho El Ocote	C_4 feeder	C_4 feeder
MPGJ-2956	<i>Dinohippus mexicanus</i>	-7.7	-6.4	Tecolotlán	C_4 feeder	C_3/C_4 feeder
MPGJ-2026	<i>Neohipparrison eurystyle</i>	-3.1	-4.8	Rancho El Ocote	C_4 feeder	C_3/C_4 feeder
MPGJ-2959	<i>Neohipparrison eurystyle</i>	-4.9	-6.4	Tecolotlán	C_4 feeder	C_3/C_4 feeder

Mesowear studies on specimens of *D. mexicanus* from the Tecolotlán basin (Barrón-Ortiz and Guzmán, 2008) found both browsers and mixed feeders.

The present isotope analysis has suggested that specimen MPGJ-2956, pertaining to *D. mexicanus* from Tecolotlán, may have fed on trees, shrubs and C_3 grasses, a diet similar to that proposed by Barrón-Ortiz and Guzmán (2008) for other Tecolotlán specimens. MPGJ-2956 was collected in the lower stratigraphic sequence of the Tecolotlán basin (Carranza-Castañeda, 2006), whereas Barrón-Ortiz and Guzmán (2008) assumed that their specimens likely came from the upper part of the sequence. Their lack of stratigraphic data precludes direct comparison of the two studies; it is uncertain whether the individuals represented by the different specimens were contemporaneous and whether the diet of this species remained constant through time. In Costa Rica, *D. mexicanus* seems to have had a generalist, not a specialized, diet based on morphological ground (Laurito and Valerio, 2010), similar to that suggested by the isotopes in specimens from Texas, Yepómera and Rancho El Ocote (MacFadden, 2008).

In the case of *N. eurystyle*, carbon isotope values for both MPGJ-2026 from Rancho El Ocote and MPGJ-2959 from the Tecolotlán basin show that these animals were C_3/C_4 mixed feeders, with in the case of MPGJ-2026, an emphasis on the intake of C_4 plants (Table 1). This pattern is different from what was found by MacFadden (2008) with specimens from Rancho El Ocote and Yepómera, which fed exclusively on C_4 plants. Mesowear analyses (Barrón-Ortiz and Guzmán, 2008) suggested that *N. eurystyle* from the Tecolotlán basin was a grazer, but the lack of stratigraphic control makes it difficult to ascertain whether the diet variation between individuals was due to different locations from the same time period, or whether the differences were due to environmental change.

Since the individuals of *N. eurystyle* analyzed here appear to have been C_3/C_4 mixed feeders like some individuals in Florida (MacFadden, 2008), both *N. eurystyle* and *D. mexicanus* from the Mexican late Hemphillian appear to have been generalists in their diet.

This plasticity in equid diets may be explained by different factors. For example, in desert areas in the southern USA and northern Mexico, C_3 plants are more abundant in winter, the rainy season, whereas C_4 plants are more abundant in summer (Ehleringer and Monson, 1993). In temperate areas in general, the two kinds of plants co-inhabit year-around, but in microhabitats that differ in temperature and humidity. Isotope analysis has shown that present-day feral horses in southern New Mexico feed on leaves and grasses that differ in proportion according to seasonal availability (Hansen, 1976; Smith *et al.*, 1998). Furthermore, factors such as soil salinity, light intensity and soil nutrients may influence the carbon isotope composition of plants (Bocherens, 2003) and hence the composition of the dental enamel of herbivores.

Competition with other herbivore species may in part explain the wide feeding spectra shown by these equids; Late-Pleistocene bison and horses at Rancho La Brea, California, fed mainly on C_3 plants, but bison more frequently preferred C_4 plants (Feranec *et al.*, 2009),

suggesting the presence of resource partitioning. The small sample size in the present study precludes further speculation.

Habitat

The carbon and oxygen isotope values of both specimens from Rancho El Ocote suggest a preference for open vegetation zones (Figure 4). In the case of the specimens from the Tecolotlán basin, *N. eurystyle* lived in grassland or savannas, whereas *D. mexicanus* preferred areas with major tree cover. In Florida, *N. eurystyle* seems to have inhabited open zones, whereas *D. mexicanus* could live in either open regions or areas with some tree cover (MacFadden *et al.*, 1999a; MacFadden, 2008). Hence, *D. mexicanus* may have been more flexible in its habitat and feeding habits according to local conditions.

Although *D. mexicanus* had a mixed feeding habit associated with environments of mixed vegetation and high humidity (Laurito and Valerio, 2010), some individuals were grazers living in herbaceous grassland and arid areas. There is a latitudinal gradient: above 35°–40°N, C_3 plants such as grasses, shrubs and trees are abundant, with diminishing C_4 grasses (MacFadden *et al.*, 1999b); this is reflected in the differences in diet between specimens from the USA and those from Mexico (MacFadden, 2008).

The results of mesowear analyses of *D. mexicanus* and *N. eurystyle* from the Tecolotlán basin suggest that this basin developed a heterogeneous vegetation of trees and grasses fostered by a seasonal climate (Barrón-Ortiz and Guzmán, 2008); dry seasons were favorable for C_4 plant growth, and wet seasons were favorable for more abundant C_3 plants. In the Santa María locality, lacustrine sediments and paleochannels indicate the presence of water bodies such as rivers,

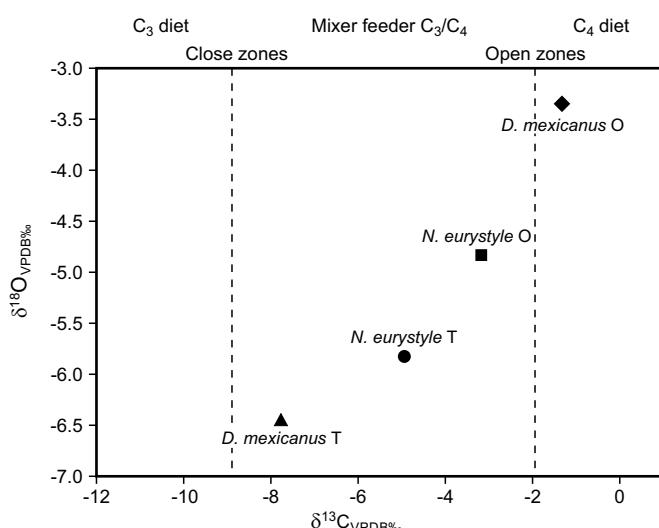


Figure 4. $\delta^{13}\text{C}$ vs $\delta^{18}\text{O}$ for equid molars from Rancho El Ocote (GTO 2A) and the Tecolotlán basin (Santa María). O: Rancho El Ocote; T: Tecolotlán.

lakes and water channels (Carranza-Castañeda, 2006). Fossil freshwater gastropods, teeth of *Crocdylus* sp., all aquatic animals, and remains of the rhinoceros *Teleoceras fossiger*, have been found there (Carranza-Castañeda *et al.*, 2005). Although *T. fossinger* has been defined as a browser, isotopic analyses of specimens from Florida showed its diet to be more generalist (MacFadden, 1998), so its presence would not necessarily indicate the existence of C₃ plants in the locality. Nevertheless, the wet environment would have been favorable to C₃ plants (Ehleringer and Monson, 1993). Abundance of C₃ plants at Santa María did not depend on the presence of a latitudinal gradient, but on specific conditions at each locality.

The conclusions above are supported by the oxygen isotope values: *D. mexicanus* from Santa María shows a δ¹⁸O value of -6.4‰, whereas from Rancho El Ocote it is -3.3‰. For *N. eurystyle* from Santa María it is -6.4‰, whereas from Rancho El Ocote it is -4.8‰. With increasing amounts of rain, oxygen isotope values become more negative (Dansgaard, 1964). Because Santa María and Rancho El Ocote are at different altitudes and latitudes, differences in δ¹⁸O values are to be expected. Additionally, the location of Santa María on the foothills of the Sierra Madre Occidental implies that it received more precipitation, creating more humidity favorable for the establishment of C₃ plants, as reflected by the lacustrine sediments (Carranza-Castañeda, 2006). Another factor, however, could explain the different oxygen isotope values. Both *D. mexicanus* and *N. eurystyle* may have migrated from place to place in search of food (MacFadden, 2008). They drank water from different feeding areas, and hence the isotopic composition could differ, as found for bison in the eastern Great Plains (Widga *et al.*, 2010).

Carbon isotope analyses on the two other equid species found at Rancho El Ocote, *A. stocki* and *N. aztecus*, showed that some of these individuals were C₃/C₄ mixed feeders, suggesting that this area was drier than at the Florida and Texas localities, and favored the presence of C₄ plants (MacFadden, 2008). However, no inference could be made regarding whether during the late Hemphillian grassland with some trees or shrubs existed.

On the basis of the diversity of fossil herbivorous species found at Rancho El Ocote or in the Tecolotlán basin, other species should be included in subsequent studies in order to draw more precise inferences regarding the vegetation type and environmental conditions that existed at each locality during the late Hemphillian.

CONCLUSIONS

At Rancho El Ocote, the diet of the *Dinohippus mexicanus* specimen was primarily based on C₄ plants, and this animal preferred open zones, such as grasslands or savannas. In contrast, the *Neohipparrison eurystyle* was a C₃/C₄ mixer feeder with an emphasis on C₄ plants, living in open zones.

In the Tecolotlán basin, the *Dinohippus mexicanus* and *Neohipparrison eurystyle* were C₃/C₄ mixed feeders and lived in areas with some tree coverage; this suggests that conditions here were more favorable for C₃ plants than at Rancho El Ocote.

These preliminary conclusions require substantiation by analyses of more specimens and from a wider range of species.

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REFERENCES

Andrade, J.L., De La Barrera, E., Reyes-Garcia, C., Ricalde, M.F., Vargas-Soto, G., Cervera, C.J., 2007, El metabolismo ácido de las crasuláceas: diversidad, fisiología ambiental y productividad: Boletín de la Sociedad Botánica de México, 87, 37-50.

Angst, D., Lécuyer, C., Amiot, R., Buffetaut, E., Fourel, E., Martinneau, F., Legendre, S., Abourachid, A., Herrel, A., 2014, Isotopic and anatomical evidence of herbivorous diet in the Early Tertiary giant bird *Gastornis*. Implications for the structure of terrestrial Paleocene ecosystems: Naturwissenschaften, 101, 313-322.

Barrick, R.E. 1998, Isotope paleobiology of vertebrates: ecology, physiology, and diagenesis, in Norris, R.D., Corfield, R.M. (eds.), Isotopic paleobiology and paleoecology 4: USA, The Paleontological Society Papers, 101-137.

Barrón-Ortiz, C.R., Guzmán, G.J.R., 2008, Hábitos alimentarios de los caballos *Dinohippus mexicanus* y *Neohipparrison eurystyle* del Henfilián Tardío de Tecolotlán, Jalisco: Investigación y Ciencia, 45, 24-29.

Bocherens, H., 2003, Isotopic biogeochemistry and the paleoecology of the mammoth steppe fauna: Deinsea, 9, 57-76.

Bravo-Cuevas, V.M., Priego-Vargas, J., Jiménez-Hidalgo, E., Barrón-Ortiz, C.R., Theodor, J.M., Castillo-Cerón, J.M., 2015, Feeding ecology of *Astrohippus stockii* from the Late Hemphillian of Central México: Stable carbon isotope analysis and dental wear patterns: Ameghiniana, 52, 502-516.

Carranza-Castañeda, O., 2006, Late Tertiary fossil localities in Central Mexico, between 19° -23° N, in Carranza-Castañeda, O., Lindsay, E.H. (eds.), Advances in late Tertiary vertebrate paleontology in Mexico and the Great American Biotic Interchange: México, Universidad Nacional Autónoma de México, Instituto de Geología y Centro de Geociencias, Publicación Especial 4, 45-60.

Carranza-Castañeda, O., Aguirre-Díaz, O., Alvarado-Ortega, J., 2005, Miocene-Pliocene paleolakes within the Central sector of the Mexican Volcanic Belt, in 65th Annual Meeting Society of Vertebrate Paleontology: Journal of Vertebrate Paleontology, 25, suplement to number 3, Abstracts of Papers, 42A-43A.

Carranza-Castañeda, O., Aranda-Gómez, J., Wang, X., Iriondo, A., 2013, The early late Hemphillian (Hh2) faunal assemblage from Juchipila Basin, State of Zacatecas, México and its biostratigraphic correlation with other Hemphillian fauna in Central México: Contributions in Science, 521, 13-49.

Cerling, T.E., 1999, Paleorecords of C₄ plants and ecosystems, in Sageand, F., Monson, R.K. (eds.), C₄ plant biology: San Diego, Ca., USA, Academic Press, 445-469.

Cerling, T.E., Harris, J.M., 1999, Carbon isotope fractionation between diet and bioapatite in ungulate mammals and implications for ecological and paleoecological studies: Oecologia, 120, 347-363.

Cerling, T.E., Ehleringer, J. R., 2000, Welcome to the C₄ world, in Gastaldo, R.A., Dimechelle, W.A. (eds.), Phanerozoic terrestrial ecosystems: USA, The Paleontological Society Papers 6, 273-286.

Cohen, K.M., Finney, S.C., Gibbard, P.L., Fan, J.-X., 2013, The ICS International Chronostratigraphic Chart: Episodes, 36, 199-204.

Coplen, T.B., 1988, Normalization of oxygen and hydrogen isotope data: Chemical Geology, 72, 293-297.

Coplen, T.B., Brand, W.A., Gehre, M., Gröning, M., Meijer Harro, A.J., Toman, B., Verkouteren, R.M., 2006, New guidelines for δ¹³C measurements: Analytical Chemistry, 78, 2439-2441.

Dansgaard, W., 1964, Stable isotopes in precipitation: Tellus, 16, 436-468.

Ehleringer, J.R., Monson, R.L., 1993, Evolutionary and ecological aspects of photosynthetic pathway variation: Annual Review of Ecology and Systematics, 24, 411-439.

Feranec, R.S., MacFadden, B., 2000, Evolution of the grazing niche in Pleistocene mammals from Florida: evidence from stable isotopes: *Palaeogeography, Palaeoclimatology, Palaeoecology*, 162, 155-169.

Feranec, R.S., Hadley, E.A., Paytan, A., 2009, Stable isotopes reveal seasonal competition for resources between Late Pleistocene Bison (*Bison*) and Horse (*Equus*) from Rancho La Brea, Southern California: *Paleogeography, Palaeoclimatology, Palaeoecology*, 271, 153-160.

Fortelius, M., Solounias, N., 2000, Functional characterization of ungulate molars using the abrasion-attrition wear gradient: a new method for reconstructing paleodiets: *American Museum Novitates*, 3301, 1-36.

Flynn, J.J., Kowallis, B.J., Nuñez, C., Carranza-Castañeda, O., Miller, W.E., Swisher III, C.C., Lindsay, E., 2005, Geochronology of Hemphillian-Blanca age strata, Guanajuato, Mexico and implications for timing of the Great American Biotic Interchange: *Journal of Geology*, 113, 287-307.

Hansen, R.M., 1976, Foods of free-roaming horses in southern New Mexico: *Journal of Range Management*, 29, 347.

Gröcker, D.R., 1997, Stable-isotopic studies on the collagenic and hydroxylapatite components of fossils: *Palaeoecological implications*: *Lethaia*, 30, 65-78.

Jiménez-Hidalgo, E., Carranza-Castañeda, O., 2010, Blancan camelids from San Miguel Allende, Guanajuato, México: *Journal of Paleontology*, 84, 51-65.

Kaiser, T.M., Solounias, N., 2003, Extending the tooth mesowear method to extinct and extant equids: *Geodiversitas*, 25, 321-345.

Koch, P.L., 1998, Isotopic reconstruction of past continental environments: *Annual Review of Earth and Planetary Science*, 26, 573-613.

Koch, P.L., 2007, Isotopic study of the biology of modern and fossil vertebrate, in Michener, R., Lajtha, K. (eds.), *Stable Isotopes in Ecology and Environmental Science*: USA, Blackwell Publishing, Boston, 99-154.

Koch, P.L., Fogel, M.L., Tuross, N., 1994, Tracing the diets of fossil animals using stable isotopes, in Lajtha, K., Michener, R.H. (eds.), *Stable isotopes in Ecology and Environmental Science*: USA, Blackwell Scientific, 63-92.

Koch, P.L., Tuross, N., Fogel, M.L., 1997, The effects of simple treatment and diagenesis on the isotopic integrity of carbonate in biogenic hydroxylapatite: *Journal Archaeological Science*, 24, 417-429.

Kohn, M.J., 1996, Predicting animal $\delta^{18}\text{O}$: accounting for diet and physiological adaptation: *Geochemical et Cosmochimica Acta*, 60, 4811-4829.

Kohn, M.J., Schoeninger, M.J., Valley, J.W., 1998, Variability in oxygen isotope compositions of herbivore teeth: reflections of seasonality or developmental physiology?: *Chemical Geology*, 152, 97-112.

Kowallis, B.J., Swisher III, C.C., Carranza-Castañeda, O., Miller, W.E., Tingey, D.G., 1998, Fission-track and single-crystal $^{40}\text{Ar}/^{39}\text{Ar}$ laser-fusion ages from volcanic ash layers in fossil-bearing Pliocene sediments in Central Mexico: *Revista Mexicana de Ciencias Geológicas*, 15(2), 157-160.

Laurito, M.C., Valerio, A.L., 2010, Los caballos fósiles de la Formación Curré, Cantón Brus. Costa Rica: Costa Rica, Museo Nacional de Costa Rica, 231 pp.

Longinelli, A., Nuti, S., 1973, Revised phosphate-water isotopic temperature scale: *Earth and Planetary Science Letters*, 19, 373-376.

MacFadden, B.J., 1998, Tale a two rhinos: Isotopic ecology, paleodiet, and niche differentiation of *Aphelops* and *Teleoceras* from Florida Neogene: *Paleobiology*, 24, 274-286.

MacFadden, B.J., 2008, Geography variations in diets of ancient populations of 5-million-year-old (early Pliocene) horses from southern North America: *Palaeogeography, Palaeoclimatology, Palaeoecology*, 266, 83-94.

MacFadden, B.J., Cerling, T.E., 1996, Mammalian herbivore communities, ancient feeding ecology, and carbon isotopes: a 10 million-year sequence from the Neogene of Florida: *Journal of Vertebrate Palaeontology*, 16, 103-115.

MacFadden, B.J., Higgins, P., 2004, Ancient ecology of 15-million-year-old browsing mammals within C3 plant communities from Panama: *Oecologia*, 140, 169-182.

MacFadden, B.J., Solounias, N., Cerling, T.E., 1999a, Ancient diets, ecology, and extinction of 5-million-year-old horses from Florida: *Science*, 283, 824-827.

MacFadden, B.J., Cerling, T.E., Harris, J.M., Prado, J.L., 1999b, Ancient latitudinal gradients of C3/C4 grasses interpreted from stable carbon isotopes of New World Pleistocene horses (*Equus*) teeth: *Global Ecology and Biogeography*, 8, 137-149.

Mahboubi, S., Bocherens, H., Scheffler, M., Benammi, M., Jeager, J.J., 2014, Was the Early Eocene proboscidean *Numidotherium hoholense* semi-aquatic or terrestrial? Evidence from stable isotopes and bone histology: *Comptes Rendus Palevol*, 13, 501-509.

Medina, E., Montes, G., Guevas, E., Rokzanidc, Z., 1986, Profiles of CO_2 concentration and $\delta^{13}\text{C}$ values in tropical rain forest of the upper Rio Negro Basin, Venezuela: *Journal of Tropical Ecology*, 2, 207-217.

Miller, W.E., Carranza-Castañeda, O., 1984, Late Cenozoic mammals from central Mexico: *Journal of Vertebrate Paleontology*, 4, 216-236.

Miller, W.E., Carranza-Castañeda, O., 1998, Late Tertiary canids from Central Mexico: *Journal of Paleontology*, 72, 546-556.

O'Leary, M.H., 1981, Carbon isotope fractionation in plants: *Phytochemistry*, 20(4), 553-567.

Pacheco-Castro, A., Carranza-Castañeda, O., 2015, Intercambio de faunas de roedores Sigmodontinos y Neotominos durante el Plioceno y Pleistoceno temprano en Norteamérica: Casos de estudio en el centro de México e inferencias paleoecológicas, in *Reunión Anual de la Unión Geofísica Mexicana 2015*: Puerto Vallarta, Jal., Mexico, *Geos*, 35 (1), 121.

Révész, K.M., Landwehr, J.M., 2002, $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ isotopic composition of CaCO_3 measured by continuous flow isotope ratio mass spectrometry: Statistical evaluation and verification by application to Devils Hole core DH-11 calcite: *Rapid Communications in Mass Spectrometry*, 16, 2012-2114.

Sánchez, B., 2005, Reconstrucción del ambiente de mamíferos extintos a partir del análisis isotópico de los restos esqueléticos, in Alcorno, P., Redondo, R., Toledo, J. (eds.), *Nuevas técnicas aplicadas al estudio de los sistemas ambientales: los isótopos estables*: Madrid, Universidad Autónoma de Madrid, 49-64.

Sanson, G.D., Kerr, S.A., Gross, K.A., 2007, Do silica phytoliths really wear mammalian teeth?: *Journal of Archaeological Science*, 34, 526-531.

Schoeninger, M.J., Kohn, M., Valley, J.W., 2000, Tooth oxygen isotope ratios as paleoclimate monitors in arid ecosystems, in Ambrose, S.H., Katzemberg, M.A. (eds.), *Biogeochemical approaches to paleodietary analysis*: New York, USA, Kluwer Academic/Plenum Publisher, 117-140.

Smith, B.N., Epstein, S., 1971, Two categories of $^{13}\text{C}/^{12}\text{C}$ ratios for higher plants: *Plant Physiology*, 47, 380-384.

Smith, C., Valdez, R., Holchek, J. L., Zwank, P. J., Cardenas, M., 1998, Diets of native and non-native ungulates in southcentral New Mexico: *Southwestern Naturalist*, 43, 163-169.

Stowe, L.G., Teeri, J.A., 1978, The geography distribution of C₄ species of the dicotyledonae in relation to climate: *The American Naturalist*, 112, 609-623.

Tedford, R.H., Albright, L.B., Barnosky, A.D.III., Ferrusquia-Villafranca, I., Hunt, R.M.Jr., Storer, J.E., Swisher, C.C.III., Voorhies, M.R., Webb, S.D., Whistler, D.P., 2004, Mammalian Biochronology of the Arikareean through Hemphillian interval (Late Oligocene through Early Pliocene Epochs), in Woodburne, M.O. (ed.), *Late Cretaceous and Cenozoic mammals of North America*: USA, Columbia University Press, 169-231.

Werner, R.A., Brand, W.A., 2001, Referencing strategies and techniques in stable isotope ratio analysis: *Rapid Communications in Mass Spectrometry*, 15, 501-519.

Widga, C., Walker, J.D., Stockli, L.S., 2010, Middle Holocene Bison diet and mobility in the eastern Great Plains (USA) based on $\delta^{13}\text{C}$, $\delta^{18}\text{O}$, and $^{87}\text{Sr}/^{86}\text{Sr}$ analyses of tooth enamel carbonate: *Quaternary Research*, 73, 449-463.

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