

# A probabilistic seismic hazard assessment of the Trans-Mexican Volcanic Belt, Mexico based on historical and instrumentally recorded seismicity

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## Resumen

La Faja Volcánica Transmexicana (FVTM) es una cadena volcánica activa que se deforma por una red de fallas extensionales intra-arco. Aunque varios sismos corticales con magnitud  $>7$  se han originado en la FVTM desde el siglo XVI, la sismicidad de fondo de esta estructura geológica es muy baja y la región es considerada, tradicionalmente, como de peligro sísmico bajo. En este estudio, presentamos un modelo probabilístico actualizado de la FVTM. El catálogo de sismicidad empleado incluye cuarenta y tres sismos registrados histórica e instrumentalmente desde 1858 hasta 2014; 5 de los cuales, son eventos de magnitud considerable ocurridos en la FVTM a lo largo del siglo XIX. Debido a la falta de una muestra estadísticamente representativa, proponemos, de manera cualitativa, que dicho catálogo es completo para magnitudes  $M \geq 4$  desde 1964 y, para magnitudes  $M \geq 6$ ; desde 1858. De igual modo, introducimos tres diferentes relaciones de magnitud-frecuencia entre eventos. La primera, es una relación convencional Gutenberg-Richter que ajusta la distribución de sismos registrados de manera instrumental. Las dos restantes, son aproximaciones semiparamétricas que integran datos históricos e instrumentados para determinar tasas de sismicidad en la región. Nuestro modelo preferido (modelo de sismicidad B) ajusta la distribución de datos históricos e instrumentados de manera independiente y

fusiona ambos ajustes en una sola curva. Por otro lado, presentamos un mapa de peligro uniforme (USH) de la FVTM para un periodo de retorno de 500 años que se obtiene considerando tres fuentes sísmicas principales: 1) Sismos de subducción con fallamiento inverso originados en la Trinchera Mesoamericana (MAT); sismos de profundidad intermedia generados dentro de la subyacente Placa de Cocos y 3) sismos de naturaleza cortical localizados en la FVTM. De acuerdo con el modelo de sismicidad B, el periodo promedio de recurrencia de sismos  $M \geq 7$  en la FVTM es, aproximadamente, 150 años. En contraste, el periodo de recurrencia calculado usando exclusivamente datos instrumentados es de 12,000 años. Los resultados de este modelo de sismicidad, que considera datos históricos e instrumentales, parecen coincidir con los periodos de retorno de sismos prehistóricos estimados para pequeños segmentos del sistema de fallas de la FVTM reportados en diversos estudios paleosismológicos. Al comparar los resultados de nuestro modelo de sismicidad predilecto, los valores de PGA obtenidos a partir de datos de sismicidad instrumentada son entre 18 y 56% menores que aquéllos predichos por el modelo de sismicidad que también considera el catálogo de sismicidad histórica.

Palabras clave: Faja Volcánica Transmexicana, catálogo de sismicidad histórica e instrumentada, modelo de sismicidad semiparamétrico, peligro sísmico.

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## Abstract

The Trans-Mexican Volcanic Belt (TMVB) is an active volcanic chain being deformed by an intra-arc extensional fault network. Although several crustal earthquakes with magnitude  $>7$  have originated in the TMVB since the 16th century, the background seismicity of this geological structure is very low and the region is usually considered of low seismic hazard. In this study, we present an updated probabilistic seismic hazard model of the TMVB. The seismicity catalog used here includes forty-three historically and instrumentally recorded earthquakes, from 1858 to 2014; five of these are large earthquakes that occurred in the TMVB during the XIXth century. Due to the lack of a statically representative sample, we propose, in a qualitative manner, the seismicity catalog is complete for  $M \geq 4$  since 1964 and for  $M \geq 6$  since 1858. Moreover, we introduce three different earthquake frequency-magnitude relations. The first one is a conventional Gutenberg Richter fit of the distribution of the instrumentally recorded earthquakes data. The other two are non-conventional, semi-parametric approaches that integrate the historical and the instrumental data to determine seismicity rates in the region. Our preferred model (seismicity model

B) fits separately the instrumental and the historical data and merge the two fits into one curve. A uniform seismic hazard (USH) of the TMVB for a return period of 500 years was calculated considering three major sources of earthquakes: 1) Subduction thrust-faulting events in the Middle American Trench (MAT); 2) Earthquakes within the subducted Cocos plate and, 3) Shallow crustal earthquakes in the TMVB. According to the seismicity model B, the average recurrence time of a  $M \geq 7$  earthquake on the TMVB is approximately 150 years. In contrast, the recurrence time estimated from the instrumental catalog is 12,000 years. The results of this seismicity model, which is based on historical and instrumental data, agrees also with the return periods of prehistoric earthquakes, estimated for short segments of the fault system in the TMVB in paleoseismological studies. When comparing the results of our preferred seismicity model, the PGA estimated using only the instrumental seismicity are 18 to 56% smaller than those predicted by the model using the historical catalog.

**Key words:** Trans-Mexican Volcanic Belt, historical and instrumental seismicity catalog, semi-parametric seismicity model, seismic hazard.

## Introduction

We present a seismic hazard assessment of the TMVB that takes into account the seismic activity observed in the region from 1858 to 2014. Although this region in central Mexico is not as seismically active as some other parts of the country, such as the Pacific coast or the Gulf of California, there is evidence of major historical crustal earthquakes ( $M > 7$ ) occurring on the TMVB in the last 500 years (García-Acosta and Suárez; 1996; Suárez and Caballero-Jiménez; 2012). As in the case of many other continental faults, the crustal faults in the TMVB responsible for these large earthquakes apparently lie dormant for several thousand years.

Although relatively infrequent, the occurrence of these shallow, crustal events in the vicinity of densely populated areas poses an important seismic hazard to the more heavily populated area of Mexico. The larger cities in the country, such as Mexico City, Guadalajara, Morelia and Jalapa lie on the TMVB. According to the latest census of the National Institute of Statistics and Geography in 2010 (<http://www.inegi.org.mx/est/contenidos/proyectos/ccpv/>), more than 43 million people (approximately

40% of the population in Mexico) live in cities and towns located on this geological province. The observations of the past 500 years indicate that crustal earthquakes of large magnitude occur throughout the TMVB, regardless of the dearth of small magnitude seismicity.

Traditionally, the seismic hazard studies conducted for cities on the TMVB are based only on the instrumental seismicity catalog. However, the background seismicity is very low and, at first glance, this suggests that the seismic hazard is moderate. The Federal Power Commission (CFE) publishes a manual for the design of civil works against earthquakes, which has become a national standard. CFE classifies the TMVB as a region of moderate hazard (<https://www.scribd.com/doc/52197523/CFE-Sismo-08>). Nevertheless, the 1912 Acambay and the 1920 Jalapa earthquakes, together with historical reports of macroseismic data, suggest that although these crustal earthquakes have long recurrence periods, they may be of relatively large magnitudes and take place in close vicinity to major cities and towns. The purpose of this paper is to integrate the large historical earthquakes of the XIX<sup>th</sup> century in order to estimate the seismic hazard in the TMVB. Our results indicate that the

hazard estimate including the historical data is much larger than previously estimated from instrumental data.

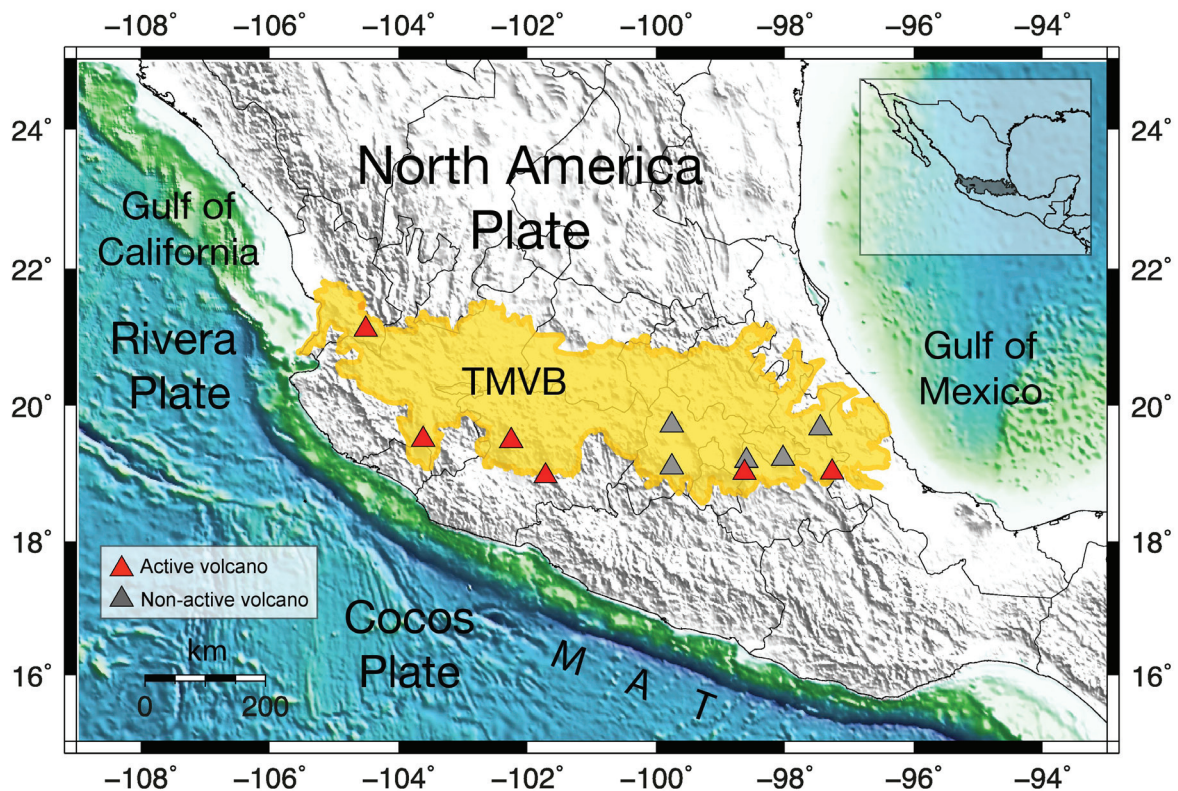
### Tectonic overview of the TMVB

The TMVB is a volcanic arc that spans Central Mexico from the Pacific coast to the Gulf of Mexico. It has a variable width that oscillates between 80 and 230 km (Figure 1). Pardo and Suárez (1993; 1995) showed that the oblique orientation of the TMVB relative to the subduction zone is due to the geometry and to the lateral changes in dip of the subducted Cocos plate beneath central Mexico (Suárez *et al.*, 1990). Pardo and Suárez (1993; 1995) speculated that to the north of the region where the seismic activity within the subducted slab ceases, the subducted Cocos plate bends sharply downwards. Recent results of the tomographic inversion of teleseismic waves observed by the dense network of temporary seismic stations of the MesoAmerican Subduction Experiment (MASE), Pérez-Campos *et al.* (2008) suggests that in central Mexico, the subducted Cocos

plate dips at an angle of  $75^\circ$  into the mantle, reaching a depth of 120 km beneath the TMVB.

The TMVB is affected mostly by extensional tectonics (e.g., Ferrari *et al.*, 2012; Suter *et al.*, 1995). Suter *et al.* (1992; 2001) attribute the extensional regime in the Morelia–Acambay fault system, in the central part of the TMVB, to isostatically compensated surface loads in response to the high elevation of the volcanic belt. Ferrari *et al.* (2012), based on the heat flow measurements of Ziagos *et al.* (1985), suggest that the origin of this extension is related to the hot and low-density mantle ( $\sim 950$  to  $1000^\circ\text{C}$ ) that underlies the Moho beneath the TMVB. This buoyant thermal effect presumably promotes the uplift and subsequent extensional deformation of the volcanic arc.

This extensional regime is characterized by seismically active, east west oriented normal faults throughout the volcanic belt. Most of these faults exhibit pronounced scarps that are clearly mapped on the surface for distances as long as 50 km (Figure 2). Many of these faults



**Figure 1.** Extent of the TMVB reported by Ferrari *et al.* (2012).



are geologically active and cut alluvium and Quaternary scoria volcanoes, indicating recent deformation. Suter *et al.* (1992) estimate an average rate slip rate of 2 mm/yr. on these extensional faults. The Quaternary vertical slip rates of these faults have a mean of 0.07 mm/yr and the bulk extension rate of the TMVB is estimated to be  $0.2 \pm 0.05$  mm/yr (Suter *et al.* 2001).

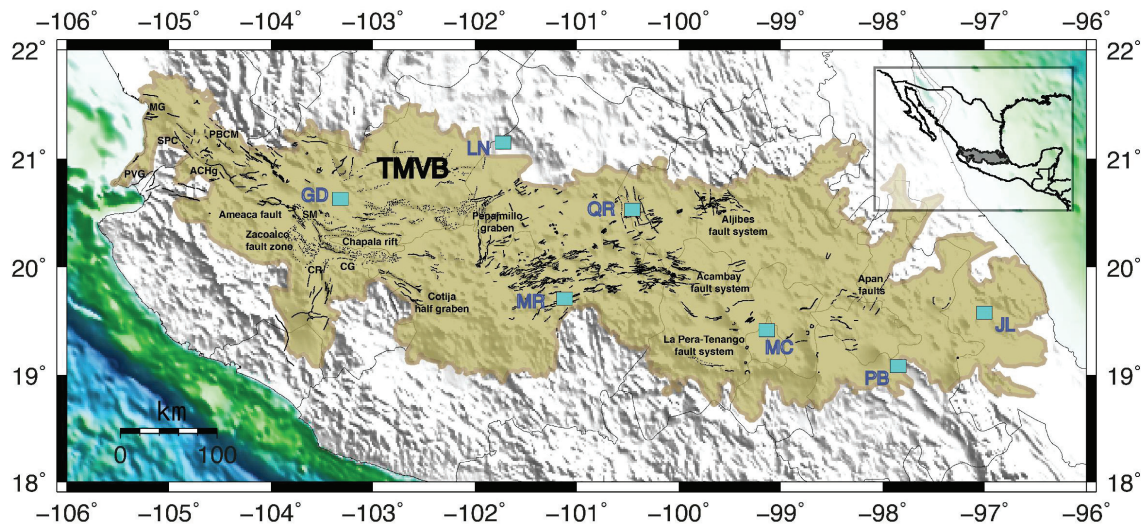
Some of these geologically active faults like the Acambay-Tixmadejé, the Venta de Bravo faults, and the faults bounding the Chapala graben, are located just a few kilometers away from densely populated cities, such as Mexico City, Guadalajara, Puebla or Morelia. Although many of these extensional crustal faults are clearly mapped on the surface, in some cases they are blind and show no surface expression. In the region of the 1920 earthquake, for example, there are no clearly identified faults mapped on the surface.

### Historical and instrumental seismicity in the TMVB

During the past 100 years, two large crustal earthquakes in the TMVB have been recorded instrumentally. On November 19, 1912, an earthquake ( $M_w$  6.9) took place near the city of Acambay rupturing one of the faults bounding the Acambay graben, about 80 km from Mexico City (Figure 3). This earthquake caused

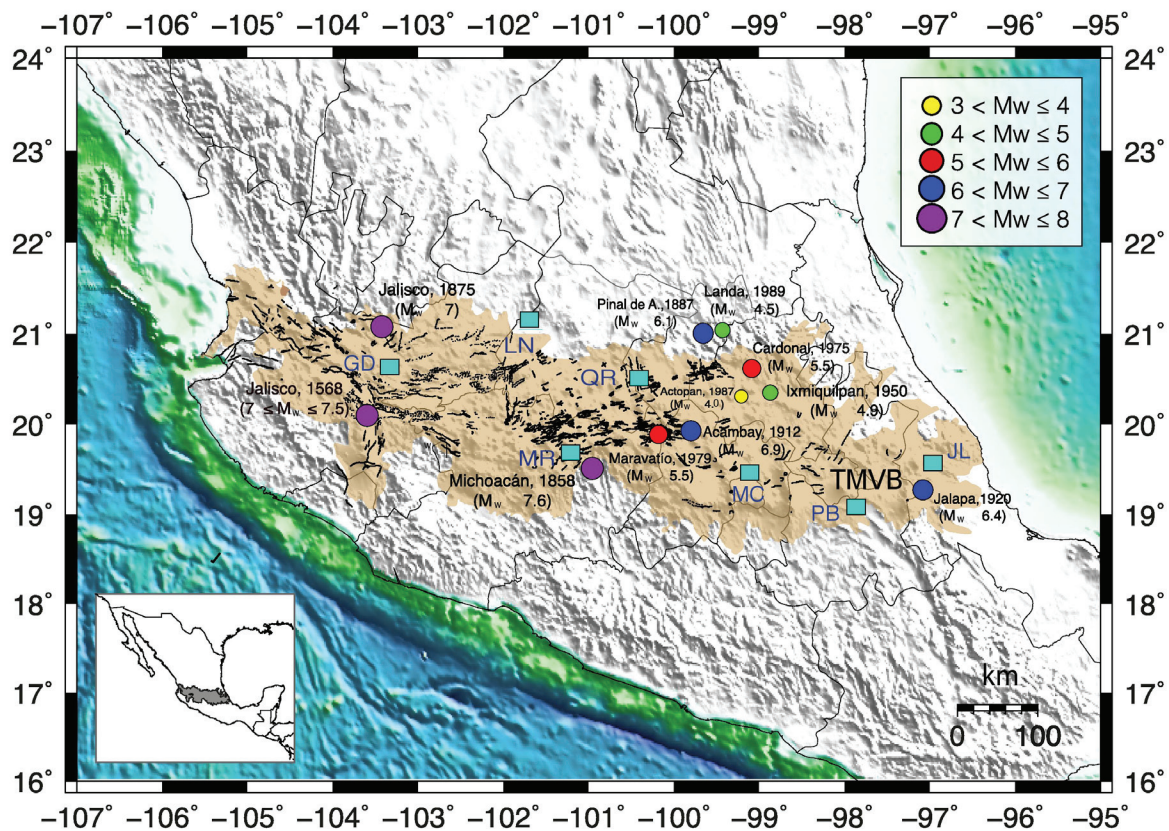
large destruction in the town of Acambay and in neighboring towns and villages (Urbina and Camacho, 1913) and caused moderate damage in Mexico City (Suter, 2015). A few years later, on January 4, 1920, an earthquake  $M_w$  6.4 took place on the eastern part of the TMVB (Figure 3). According to Suárez (1992), this event is one of the more deadly earthquake ever recorded in Mexico, just after the large 1985 Michoacán earthquake ( $M_w$  8.1). More than 1,500 fatalities were estimated. Most of them drowned or buried under the debris flows caused by the landslides on the steep cliffs bounding the course of the Pescados river (Comisiones del Instituto Geológico Mexicano, 1922). The most recent moderate-sized earthquake in the TMVB took place on February 22, 1979. The Maravatio earthquake ( $m_b$  5.3) occurred near the western end of the Acambay fault and caused only slight damage (Astiz, 1980).

The recurrence period of some segments of these crustal faults were estimated on the basis of paleoseismological studies in the vicinity of the Acambay graben. Langridge *et al.* (2000) conducted a study on the fault trace of the 1912 Acambay earthquake. Their results show the presence of three previous ruptures with an approximate recurrence time of 3,600 years. Similar paleo seismological studies conducted on the neighboring faults, on the southern boundary of the Acambay graben,



**Figure 2.** Distribution of the faults with Quaternary activity mapped on the TMVB. PVG represents the Puerto Vallarta graben; MG Mecatán Graben; SPC the San Pedro-Ceboruco Graben; ACH the Amatlán de Cañas Half-graben; PBCM the Plan de Barrancas-Cinco Minas Graben; SM the San Marco fault; CR the Colima Rift; CG the Citlala Graben. The major cities are abbreviated as: GD: Guadalajara; LN: León; MR: Morelia; QR: Querétaro; MC: Mexico City; PB: Puebla; and JL: Jalapa.





**Figure 3.** Spatial and temporal distribution of the more important historical and instrumental seismicity within the TMVB since the XVI<sup>th</sup> century.

indicate rupture intervals ranging from 600 to 10,000 years, depending on whether each fault breaks as a single unit or in separate segments (Langridge *et al.*, 2013; Ortuño *et al.*, 2015). Thus, based on the paleo seismological data, a significant earthquake may be expected in this region of central Mexico every 300 to 600 years (Zúñiga *et al.*, 2012).

Prior to the instrumental period, several large earthquakes have been identified as crustal events in the TMVB, based on the damage and felt reports. The earliest reference is the December 27, 1568, earthquake (Figure 3). Fortunately, for an earthquake that occurred so early in the written historical record of Mexico, there are detailed descriptions of the damage and of the substantial and extended ground deformation caused by this earthquake (Suárez *et al.*, 1994). Practically all of the newly built Franciscan churches in this area, located to the southwest of the city of Guadalajara, were destroyed. Based on the area of large seismic intensity, Suárez *et al.*, (1994)

suggest a magnitude of  $M_w$  7.0. Later, Suter (2015) suggested a magnitude  $M_w$  7.2 for this earthquake interpreting the ground rupture reported by historical sources as evidence of the fault trace. There are contradictions in the historical reports regarding the year of this earthquake. Suter (2015) suggests that this earthquake actually took place in 1567 and not in 1568.

Based on the inversion scheme of seismic intensity data proposed by Bakun and Wentworth (1997), Suárez and Caballero-Jiménez (2012) and Suárez *et al.* (2016) determined the approximate epicentral location and the magnitude of three large earthquakes in the TMVB that occurred during the XIX<sup>th</sup> century. The Santa Juliana earthquake on June 19, 1858, ( $M$  7.6), the San Cristobal event of February 11, 1875, ( $M$  7.0) and the Pinal de Amoles earthquake of November 26, 1887, ( $M$  6.1) have numerous and detailed historical reports (García Acosta and Suárez, 1996). The number and quality of the macroseismic

data available was sufficient to allow for a quantitative determination of the source parameters using the inversion scheme of the observed Modified Mercalli Intensities (MMI) as a function of distance (Figure 3). Singh *et al.* (1996) interpreted the Santa Juliana earthquake as an in-slab event. Suárez and Caballero-Jiménez (2012) and Suárez *et al.* (2016) provide evidence that this event is a crustal event.

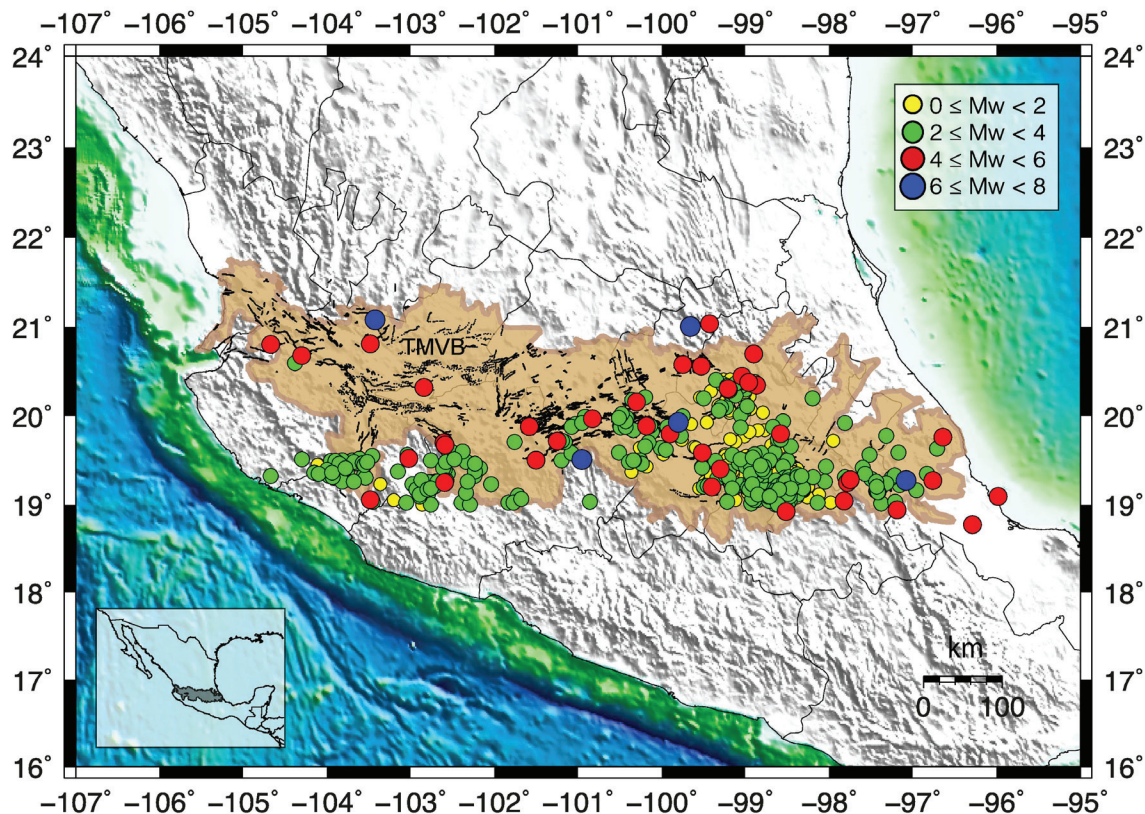
### Probabilistic seismic hazard assessment

The seismic hazard analysis presented here was developed following the methodology proposed by Cornell (1968) and Esteva (1968). To this end, we compiled a robust seismicity catalog identifying the location of the seismic sources of interest that allows us to compute the mean recurrence periods of earthquakes in the region, as a function of magnitude. The seismic hazard estimations were obtained using the computer program *CRISIS 2015*, an updated version of *CRISIS 2007* (Ordaz *et al.*, 2007; <https://ecapra.org/crisis-2007>).

### Seismicity catalog and its level of completeness

The seismicity catalog used was based on an exhaustive search of the seismic databases. Thus the data were culled from the catalogs of the International Seismological Centre (ISC), the U.S. Geological Survey (USGS), the Global Centroid Moment Tensor Catalog (CMT), and of the Servicio Sismológico Nacional (SSN) were consulted. Duplications and aftershocks were avoided.

A total of 567 historical and instrumental earthquakes were initially selected. A maximum depth cutoff of 33 km was defined to guarantee that the earthquakes occurred in the crust (Figure 4). Due to the fact that only moderate and large earthquakes contribute to large ground accelerations, only  $M \geq 4$  earthquakes were taken into account in this study. Furthermore, due to the dearth of seismic stations in the region, events  $M < 4$  are poorly located and there is a great uncertainty in their focal depth. In fact, many of the small earthquakes presumed to be in the TMVB crust are probably deeper, in-slab events within the subducted Cocos plate (green circles on Figure 4).



**Figure 4.** Epicentral locations of earthquakes observed on the TMVB from 1858 to 2012. These events represent the selected seismic catalog that constitutes the basis of the seismic hazard analysis presented here.



A selected dataset of 43 events satisfied the selection criteria and constitute the basis of this analysis. The construction and operation of modern seismic networks, such as the national network of the SSN and the Worldwide Standard Seismic Network (WWSSN) substantially improved the detection capability of small earthquakes. Since 1964 the number of located small earthquakes has increased dramatically. Zúñiga *et al.* (2016) propose that the seismicity catalog of Mexico is complete for events approximately  $m_b > 3.6$  since 1964. This observation is based on a completeness analysis of the Mexican catalog. The temporal changes and threshold levels observed in the Mexican seismicity catalog are similar to the variations observed worldwide (e.g., Zúñiga and Wyss 1995; Zúñiga *et al.* 2005; Michael 2014). In a conservative manner, we consider here that the seismicity catalog of the TMVB is complete for  $M \geq 4$  since 1964.

On the other hand, the population on the TMVB has been relatively dense since the XIX<sup>th</sup> century. As the historical catalog demonstrates, it is very likely that the damage and effects of moderate sized earthquakes would have been reported in the historical accounts. Thus we qualitatively suggest that the catalog is complete for earthquakes  $M \geq 6$  since the second part of the XIX<sup>th</sup> century.

#### Seismic source characterization

Seismic faults on the TMVB are mapped throughout this geological province (Figures 2 and 3). Although some faults indicate recent deformation, there is insufficient geological information to state categorically that all mapped faults have been active in the Quaternary. On the other hand, it is important to point out that large crustal earthquakes have taken place in regions of the TMVB where there are no surface mapped faults. The apparent absence of faults in regions where important earthquakes have occurred may be due to insufficient mapping or to the fact that some faults are blind and do not necessarily outcrop at the surface. The more relevant example of earthquakes occurring where there are no geologically mapped faults is the 1920 Jalapa earthquake ( $M_w$  6.4).

Considering that large historical and instrumental earthquakes are almost homogeneously distributed in the TMVB (Figures 2 and 3), we assume here that seismically capable crustal faults in the TMVB are also homogeneously distributed. This assumption means that the probability to

expect an earthquake higher than a certain magnitude (in this case  $M_{max} = 7.6 \pm 0.3$ , the maximum magnitude earthquake observed in the region) is the same for every location of the TMVB (Suárez and Caballero-Jiménez, 2012; Suárez *et al.*, 2016). Admittedly, this assumption may lead to overestimate the peak ground accelerations expected in the zone, as there is not enough evidence to confirm that all of these faults exist or remain active in the Quaternary. Thus we consider this assumption of uniformly distributed seismic hazard of the TMVB, as the more conservative assessment that can be made on the basis of the existing information.

#### Seismicity models

The more frequently used statistical model to represent the rate of occurrence of earthquakes is the Gutenberg-Richter relation. This relation establishes that the number of earthquakes is a function of the magnitude. This frequency magnitude relation is expressed as:

$$\log_{10} N = a - bM \quad (1)$$

This relation between magnitude  $M$  and the logarithm of the number of earthquakes  $N$  is parameterized by the intercept of  $a$  and the slope  $b$  of the linear fit (Ishimoto and Ida, 1939; Gutenberg and Richter, 1944). A more conservative expression of earthquake occurrence is obtained by associating a maximum magnitude,  $M_{max}$ , to the distribution (Aki, 1965). The modified seismicity model is expressed as:

$$\lambda(M) = \lambda_c (e^{-\beta M} - e^{-\beta M_{max}}) / (e^{-\beta M_c} - e^{-\beta M_{max}}) \quad (2)$$

where,  $\lambda_c = N/A$  is the exceedance rate of the minimum magnitude earthquake ( $M_c$ ) in the catalog.  $\beta = 1/(M_p - M_c) = b \ln(10)$ .  $N$  is the number of earthquakes with magnitudes  $M \geq M_c$ , the parameter  $A$  is the duration of the seismicity catalog and  $M_p$  is the average magnitude of all the events (Aki, 1965).

Three seismicity models were constructed based on the compiled seismicity catalog. The first one, Model A, is a traditional parametric model following the modified Gutenberg-Richter relation and using only the instrumental catalog (green curve on Figure 5). The exceedance rate for a cut-off magnitude  $\lambda_c$  and the  $\beta$  value were computed using the maximum likelihood method proposed by Aki (1965). The resulting seismicity parameters for this model are  $\lambda_c = 3.333$  and  $\beta = 0.750$ . These values are similar

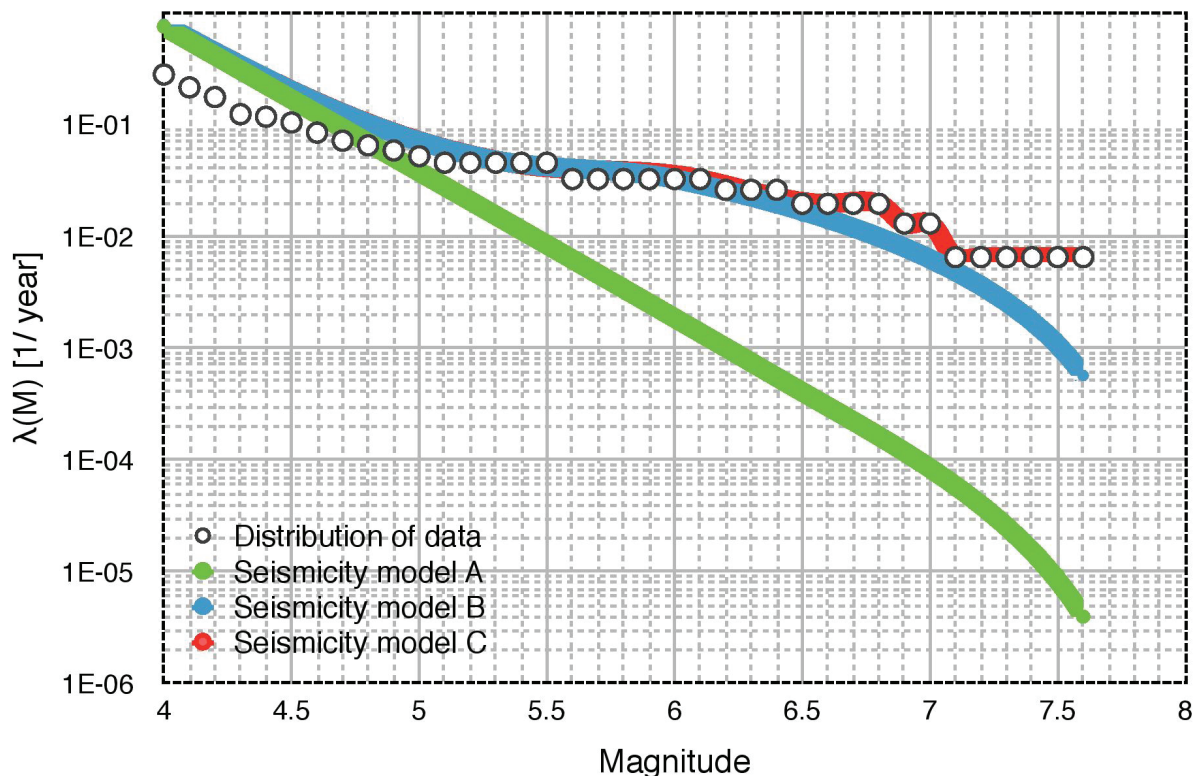


to those reported by Ordaz *et al.*, (2012), which are the current values used to quantify the seismic hazard of the TMVB.

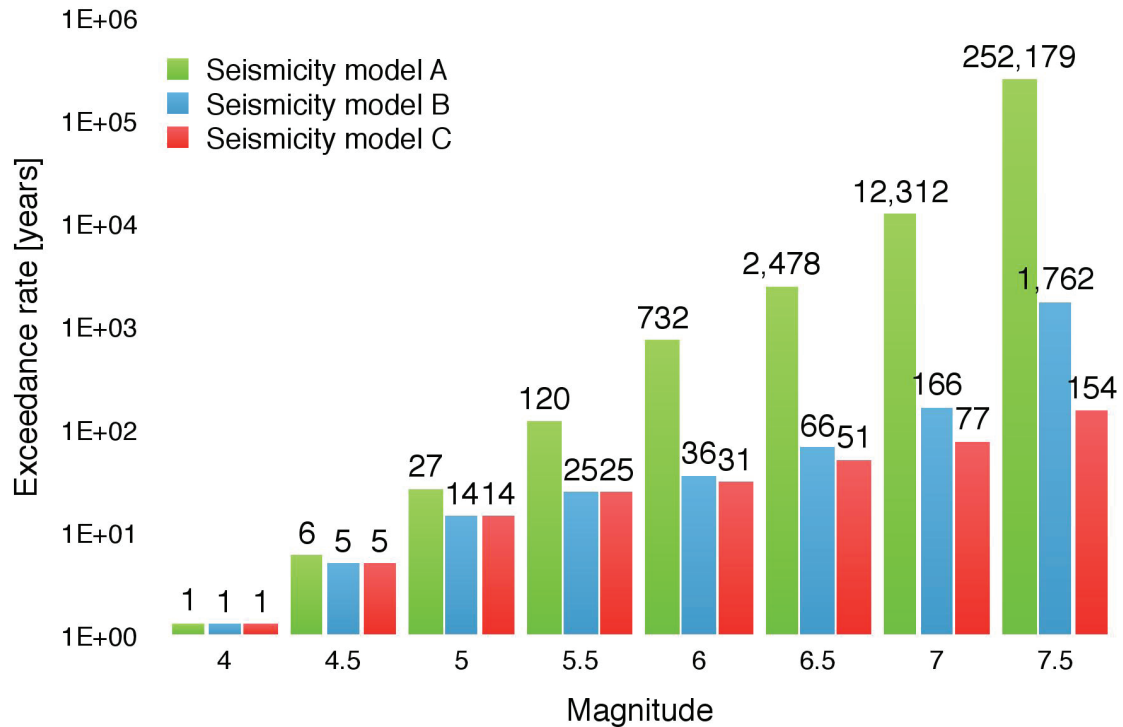
Two other non-parametric models are proposed that integrate the historical and the instrumental seismic data (Figure 5). The seismicity Model B (blue line on Figure 5) represents an effort to describe the seismic activity of the TMVB in a semi-parametric manner, using both the instrumental and the historical earthquakes of the catalog. In this approach, two different Gutenberg-Richter distributions are estimated: one of the curves corresponds to the instrumental data and the other to the historical seismicity. Thus considering the assumption that the seismicity catalog is complete for magnitudes  $M \geq 4$  since 1964 and for  $M \geq 6$  since 1858, we estimate one Gutenberg-Richter linear function for earthquakes  $4 \leq M < 6$  earthquakes, and a separate one for earthquakes  $M \geq 6$  events. As we have not observed any earthquake  $M \geq 6$  on

the TVMB since 1964, we propose to fuse both fits into one curve by adding the exceedance rate of instrumental earthquakes to historical events with  $M > 6$ . The resulting seismic parameters are:  $\lambda_c = 0.033$  and  $\beta = 1.282$

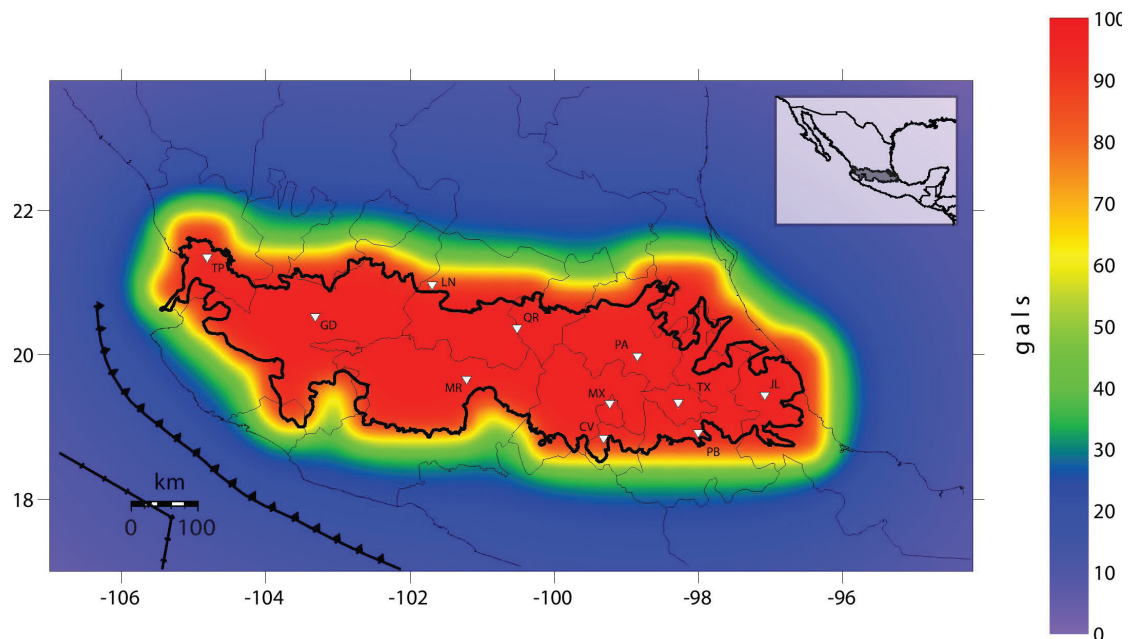
Finally, a third model (Model C and red curve on Figure 5) results from fusing the exceedance rates obtained from the fit of instrumental data ( $4 \leq M < 6$ ) with the exceedance rates of  $M \geq 6$  earthquakes observed from the catalog. This model represents the upper bound of peak ground accelerations PGA that would be expected in this geological province. The mean recurrence times of earthquakes of a certain magnitude that may occur anywhere in the TMVB, according to the three seismicity models, are shown on Figure 6. Notice that for earthquakes  $M_w > 5.5$ , the recurrence time between Model A, using only instrumental data, is substantially shorter than for the two other models, which include historical data.



**Figure 5.** The exceedance rates of earthquakes expected in the TMVB are shown as a function of magnitude. Model A (red curve) is obtained from the fit of instrumental data according to a conventional Gutenberg-Richter distribution and the maximum likelihood method. Model B (blue curve) results from merging the historical and instrumental data. This semi parametric approach appears to be a conservative description of the exceedance rates of earthquakes observed in the TMVB since the 16th century. Finally, seismicity Model C (green curve) is obtained by fusing the exceedance rates reproduced by the fit of instrumental data ( $4 \leq M < 6$ ) with the exceedance rates of earthquakes  $M > 6$  observed from the seismic catalogs.



**Figure 6.** Mean return periods of crustal earthquakes expected on the TMVB as a function of magnitude. These estimations are computed using the three different seismicity models discussed. For earthquakes  $M > 5.5$ , the results from the traditional model that uses only instrumental data (green bars) begin to differ significantly with the results predicted by the semi parametric models (bars in blue and red), which also consider historical seismicity.



**Figure 7.** Seismic hazard map of the TMVB for a return period of 500 years. Peak ground accelerations are estimated thinking the TMVB as a whole and assuming the semi-parametric seismicity model B as the preferred model describing seismicity rates in the volcanic arc. Such estimations might represent upper boundaries of PGA, however, they seem to fit the few observations available to date.

### Seismic hazard estimations of the TMVB

We propose that model B is a conservative estimation of the seismicity in the TMVB. Using this model, we estimate the seismic hazard of the TMVB and a uniform seismic hazard (USH) map of the central portion of Mexico. In order to estimate the USH of the TMVB, three types of earthquake sources are considered: 1) shallow crustal seismicity occurring on the TMVB; 2) in-slab earthquakes occurring within the subducted Cocos plate and; 3) subduction zone earthquakes along the Middle American Trench (MAT).

#### Attenuation models

Due to the lack of strong motion records of crustal earthquakes in the TMVB, it was necessary to use Ground Motion Prediction Equations (GMPE) from other regions of the world with similar geological characteristics. We use the ground motion prediction equations reported by Abrahamson *et al.*, (2014), which describe the attenuation of the spectral response values in the horizontal and vertical components, for crustal earthquakes in tectonically active regions.

In the case of in-slab earthquakes, the parameters reported by Ordaz *et al.*, (2012) and the GMPE's proposed by García *et al.*, (2005) were utilized. On the other hand, the seismicity rate in the Middle American Trench was determined using a traditional Gutenberg- Richter distribution. This was obtained compiling information of 288 thrust earthquakes reported in the Global Centroid Moment Tensor (CMT) database (Dziewonsky *et al.*, 1981). Earthquakes from this catalog were selected from January 1<sup>st</sup>, 1976 to December 31, 2014, within the geographic coordinates  $15^{\circ}\text{N} \geq \text{latitude} \leq 24^{\circ}\text{N}$  latitudes and  $106^{\circ}\text{W} \leq \text{longitude} \leq 92^{\circ}\text{W}$ . All the events have magnitudes  $M_w \geq 4.7$  and focal depths of less than 40 km. In this case, the GMPE's estimated by Arroyo and Ordaz, (2010). for subduction zone earthquakes were utilized.

In computing ground motion expected for the return periods of 100, 500, 2500 and 10,000 years, we assumed a Poisson model to describe the probabilities of an earthquake occurrence. In other words, the seismicity rates on each seismic source are time-independent. The resulting seismicity parameters for the various seismic sources are summarized on Table 1.

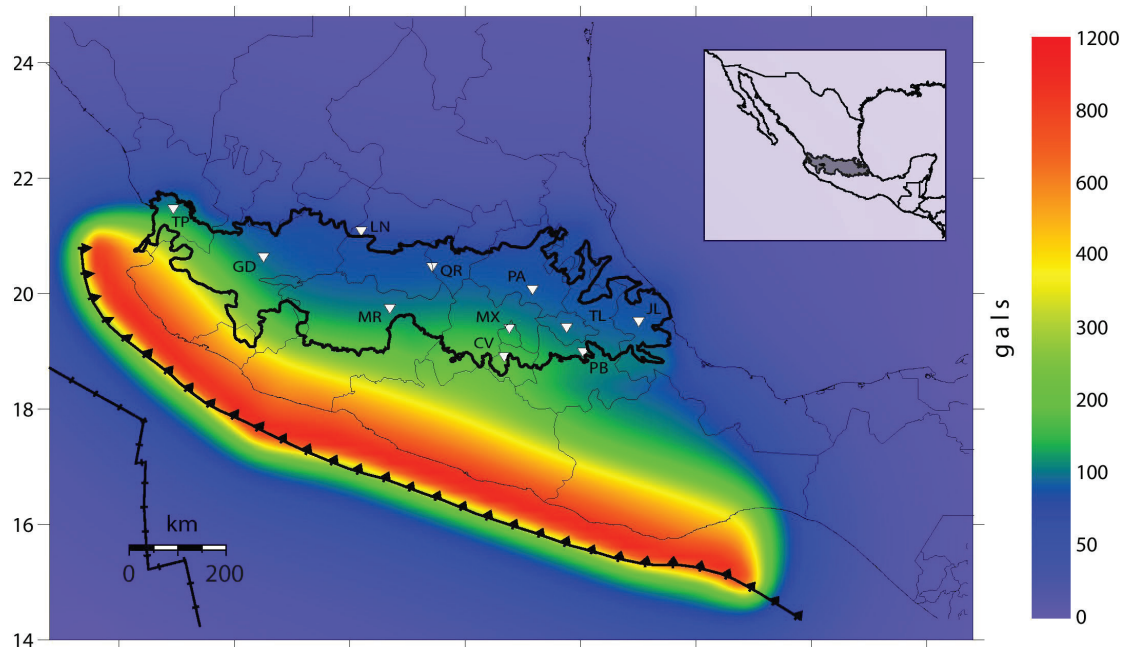


Figure 8. Uniform hazard map of central Mexico for a return period of 500 years. This model is built considering subduction thrust-faulting earthquakes occurring in the Middle American Trench, normal-faulting events within the subducted Cocos plate and shallow crustal earthquakes in the TMVB. The seismicity parameters used to create this model are reported in Table 1.



**Table 1.** Seismicity parameters used in this study to assess the seismic hazard of the TMVB.

Type of faulting	$\lambda$	$\beta$	$\delta\beta$	$M_{max}$	$M_c$	GMPE
Crustal instrumental earthquakes	0.750	3.333	0.17	7.6	4.0	Abrahamson <i>et al.</i> , (2014)
Crustal historical earthquakes	0.033	1.282	0.45	7.6	6.0	Abrahamson <i>et al.</i> , (2014)
In-slab earthquakes (Central Mexico)	9.063	2.590	0.04	8.1	4.0	García <i>et al.</i> , (2005)
In-slab earthquakes (Western Mexico)	6.452	2.078	0.04	8.1	4.0	García <i>et al.</i> , (2005)
Subduction Earthquakes	7.385	1.235	0.06	8.1	4.7	Arroyo <i>et al.</i> , (2010)

### Discussion of results and conclusions

In this study, we propose that a non-traditional, semi-parametric seismicity model, which incorporates instrumental and historical data, represents better the seismicity of the TMVB and is better suited to assess the seismic hazard. For example, according to the seismicity Model B, the average recurrence time of a  $M \geq 7$  earthquake on the TMVB is approximately 150 years. In contrast, the mean recurrence time estimated only on the instrumental catalog (Model A) is 12,000 years. The historical earthquake record of the past 160 years indicates that two earthquakes  $M \geq 7$  have taken place on the TMVB. The mean recurrence time predicted by Model B of earthquakes  $M \geq 6.5$  is of 66 years. The seismicity catalog shows that there are four events in this magnitude range in the last 160 years. Again, this is in good agreement with the predictions made by Model B and clearly shows that using only instrumental data, recurrence times of moderate and large earthquakes in the TMVB are considerably underestimated.

If we also consider the data available from paleo seismological studies, a similar picture emerges (Table 2). Studies performed in segments of crustal faults located on the TMVB indicate that the mean return periods of large earthquakes ( $M_w > 7.0$ ) are in the order of 3,600 to 12,500 years (Table 2). The return period predicted by instrumental data is 2,600 years. However, it should be stressed that the return periods estimated from paleo seismological studies refer only to short segments of the fault system in the TMVB. Clearly, considering the whole system of crustal faults observed in the TMVB, the return period predicted by Model A would be an underestimation.

Due to the epistemic uncertainties stemming from the limited data available, it was necessary to generate scenarios representing upper and lower bounds of the seismic hazard in the region. These models estimate PGA's on hard soil of the TMVB for the more densely populated locations located on this geological province based on the uniform seismic hazard assessment presented above (Tables 3 and 4).

**Table 2.** Paleoseismological studies carried out in segments of crustal faults located on the TMVB.

Reference	Fault segment	Average observation window	Number of Earthquakes	Mean Return Period (yrs)	Magnitude $M_w$
Langridge <i>et al.</i> , (2000)	Acambay	~ 11,570	3	3,600	6.9
Langridge <i>et al.</i> , (2013)	Pastores	~ 35,000	3	(10,000 - 15,000)	7.0
Suñe-Pujol <i>et al.</i> , (2014)	Temascalcingo	~ 27,000	3	9,000	6.4
Ortuño <i>et al.</i> , (2015)	Pastores (west segment)	~ 4,000	5	(1,000 - 2,600)	Variable rupture

**Table 3.** Peak ground accelerations estimated for basin fill of the TMVB according to three different models describing its seismicity rate.

Return period	Seismicity model A	Seismicity model B	Seismicity model C
100 years	14 gals	28 gals	35 gals
500 years	41 gals	102 gals	133 gals
2,500 years	98 gals	260 gals	347 gals
10,000 years	181 gals	481 gals	623 gals

**Table 4.** Uniform seismic hazard estimations expected in Mexico City, Guadalajara, Morelia and Jalapa, according to three seismicity models of the TMVB, for different return periods.

## Mexico City (19° 26' N, 99° 8' W)

Return period	Seismicity model A	Seismicity model B	Seismicity model C
100 years	45 gals	63 gals	70 gals
500 years	130 gals	164 gals	190 gals
2,500 years	271 gals	348 gals	411 gals
10,000 years	425 gals	570 gals	698 gals

## Guadalajara (20° 40' N, 103° 21' W)

Return period	Seismicity model A	Seismicity model B	Seismicity model C
100 years	38 gals	49 gals	63 gals
500 years	91 gals	134 gals	166 gals
2,500 years	182 gals	299 gals	374 gals
10,000 years	303 gals	500 gals	668 gals

## Morelia (19° 46' 06" N, 101° 11' 22" W)

Return period	Seismicity model A	Seismicity model B	Seismicity model C
100 years	45 gals	59 gals	66 gals
500 years	112 gals	149 gals	176 gals
2,500 years	221 gals	319 gals	386 gals
10,000 years	353 gals	523 gals	674 gals

## Jalapa (19° 32' 24" N, 96° 55' 59" W)

Return period	Seismicity model A	Seismicity model B	Seismicity model C
100 years	27 gals	34 gals	38 gals
500 years	65 gals	109 gals	136 gals
2,500 years	141 gals	275 gals	355 gals
10,000 years	241 gals	484 gals	656 gals

**Table 5.** Comparison of the peak ground accelerations expected in Mexico City, Guadalajara, Morelia and Jalapa, for a return period of 500 years, according to Ordaz (2004) and this article.

Seismicity Model	Mexico City	Guadalajara	Morelia	Jalapa
Ordaz (2004)	~ 125 gals	~ 75 gals	~ 100 gals	~ 90 gals
This study, model A	130 gals	91 gals	112 gals	65 gals
This study, model B	164 gals	134 gals	149 gals	109 gals
This study, model C	190 gals	166 gals	176 gals	136 gals

It is worth to compare the results obtained here for the more important cities on the TMVB with the values of PGA proposed by Ordaz (2004) for a return period of 500 years (Table 5). The PGA values estimated by Ordaz (2004) are almost identical to those estimated from Model A in this study. This is not surprising, considering that both use a very similar instrumental catalog. However, when comparing with the results of Model B, the PGA's estimated using only the instrumental seismicity are 18 to 56% smaller than those predicted in the model using the historical catalog. The largest differences are in the cities of Morelia and Guadalajara, due to the proximity of large historical earthquakes (Table 5).

In conclusion, our results suggest strongly that the estimations of PGA's may be severely underestimated using the limited instrumental catalog. Important differences exist in the uniform seismic hazard estimations with the incorporation of historical seismicity into the seismicity catalog. The seismic hazard determined with the use of the historical data, albeit only available for a short period, agree better with the seismic record and with the paleo seismological observations available. Clearly, this is a first approach to the important issue of estimating seismic hazard in this highly populated region of Mexico. It is evident that as more seismic, geological and geodetic data become available, hazard estimates should be improved. In particular, future work should consider including the local site response of the larger cities on the TMVB, particularly in Mexico City, in order for these results to be incorporated in the building codes of central Mexico.

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### References

- Abrahamson N.A., Silva W.J., Kamai R., 2014, Summary of the ASK14 ground motion relation for active crustal regions. *Earthquake Spectra*, 30, 3, pp.1025-1055.
- Aki K., 1965, 17. Maximum likelihood estimate of  $b$  in the formula  $\log N = a - bM$  and its confidence limits.
- Arroyo D., Ordaz M., 2010, Multivariate Bayesian regression analysis applied to ground-motion prediction equations, Part 2: Numerical example with actual data. *Bulletin of the Seismological Society of America*, 100, 4, 1568-1577.
- Astiz L.M., 1980, Sismicidad en Acambay, Estado de México. El temblor del 22 de febrero de 1979. *Professional thesis, Mexico City, Universidad Nacional Autónoma de México, Facultad de Ingeniería*.
- Bakun W.U., Wentworth C.M., 1997, Estimating earthquake location and magnitude from seismic intensity data. *Bulletin of the Seismological Society of America*, 87, 6, 1502-1521.
- Comisiones del Instituto Geológico de México, 1922, Memoria relativa al terremoto mexicano del 3 de enero de 1920: Instituto Geológico de México, Boletín 38, 106 p, <http://bcct.unam.mx/bogeolpdf/geo38/>.
- Cornell C.A., 1968, Engineering seismic risk analysis. *Bulletin of the Seismological Society of America*, 58, 5, 1583-1606.



- Dziewonski A.M., T.-A. Chou and J. H. Woodhouse, (1981). Determination of earthquake source parameters from waveform data for studies of global and regional seismicity, *J. Geophys. Res.*, 86, 2825-2852, doi:10.1029/JB086iB04p02825.
- Esteve, L. (1968). *Bases para la formulación de decisiones de diseño sísmico*. Instituto de Ingeniería, Universidad Nacional Autónoma de México.
- Ferrari, L., Orozco-Esquivel, T., Manea, V., and Manea, M. (2012). The dynamic history of the Trans-Mexican Volcanic Belt and the Mexico subduction zone. *Tectonophysics*, 522, 122-149.
- García-Acosta, V., and Suárez, G. (1996). Los sismos en la historia de México, tomo I. *Universidad Nacional Autónoma de México/ Centro de Investigaciones y Estudios Superiores en Antropología Social/Fondo de Cultura Económica*, 719.
- García, D., Singh, S. K., Herráiz, M., Ordaz, M., and Pacheco, J. F. (2005). Inslab earthquakes of central Mexico: peak ground-motion parameters and response spectra. *Bulletin of the Seismological Society of America*, 95(6), 2272-2282.
- Gutenberg R. and Richter C., 1944, Frequency of earthquakes in California, *Bulletin of the Seismological Society of America*, 34, pp. 185-188.
- Ishimoto M., Ida K., 1939, Observations sur les seisms enregistres par le microseismograph construit dernièrement, *Bull. Earthqu. Res. Int.*, 17, 443-478.
- Langridge, R. M., Weldon, R. J., Moya, J. C., & Suárez, G. (2000). Paleoseismology of the 1912 Acambay earthquake and the Acambay-Tixmadejé fault, Trans-Mexican volcanic belt. *Journal of Geophysical Research: Solid Earth*, 105(B2), 3019-3037.
- Langridge, R. M., Persaud, M., Zúñiga, F. R., Díaz, G. D. J. A., Pérez, P. V., and Lacan, P. (2013). Preliminary paleoseismic results from the Pastores fault and its role in the seismic hazard of the Acambay graben, Trans-Mexican Volcanic Belt, Mexico. *Revista Mexicana de Ciencias Geológicas*, 30(3), 463-481.
- Michael, A., (2014). How Complete is the ISC-GEM Global Earthquake Catalog?, *Bull. of the Seism. Soc. Am.*, 104, 1829-1837, doi: 10.1785/0120130227
- Ordaz, M. (2004). Mapa de peligro sísmico de México con fines de calificación de equipo eléctrico, Instituto de Ingeniería, UNAM, Mexico.
- Ordaz, M., A. Aguilar and J. Arboleda (2007). Program for Computing Seismic Hazard: CRISIS 2007 v1.1, Instituto de Ingeniería, UNAM, México.
- Ordaz, M., D. Arroyo, S.K. Singh and G. Suárez (2012). Estudios para la reevaluación del riesgo sísmico para el sitio de la Central Nucleoeléctrica Laguna Verde (CNLV), Instituto de Ingeniería, UNAM, México.
- Ortuño, M., F. R. Zúñiga, G. J. Aguirre-Díaz, D. Carreon-Freyre, M. Cerca, and M. Roverato (2015). Holocene paleo-earthquakes recorded at the transfer of two major faults: The Pastores and Venta de Bravo faults (Trans Mexican Volcanic Belt), *Geosphere*, v 11, 2, 160-184).
- Pardo, M., and Suárez, G. (1993). Steep subduction geometry of the Rivera plate beneath the Jalisco block in western Mexico. *Geophysical Research Letters*, 20(21), 2391-2394.
- Pardo, M., and Suárez, G. (1995). Shape of the subducted Rivera and Cocos plates in southern Mexico: Seismic and tectonic implications. *Journal of Geophysical Research: Solid Earth* (1978-2012), 100(B7), 12357-12373.
- Pérez-Campos, X., Kim, Y., Husker, A., Davis, P. M., Clayton, R. W., Iglesias, A. and Gurnis, M. (2008). Horizontal subduction and truncation of the Cocos Plate beneath central Mexico. *Geophysical Research Letters*, 35(18).
- Singh, S.K., Ordaz, M. and Pérez-Rocha, L.E., 1996. The great Mexican earthquake of 19 June 1858: Expected ground motions and damage in Mexico City from a similar future event. *Bulletin of the Seismological Society of America*, 86, 6, pp.1655-1666.
- Suárez, G., Monfret, T., Wittlinger, G., and David, C. (1990). Geometry of subduction and depth of the seismogenic zone in the Guerrero gap, Mexico. *Nature*, 345(6273), 336-338.
- Suárez, G. (1992). El sismo de Jalapa del 3 de Enero de 1920. *Rev. Mex. Ing. Sísm.*, 42, 3-15.
- Suárez, G., García-Acosta, V., and Gaulon, R. (1994). Active crustal deformation in the Jalisco block, Mexico: evidence for a great

- historical earthquake in the 16th century. *Tectonophysics*, 234(1), 117-127.
- Suárez, G. and Caballero-Jiménez, G.V. (2012). Quantitative Evaluation of Historical Earthquakes on the Mexican Volcanic Belt, 3<sup>rd</sup> INQUA-IGPC-567, *International Workshop on Active Tectonics, Paleoseismology and Archaeoseismology, Morelia, Mexico*.
- Suárez, G., G.V. Caballero-Jiménez, D.A.-Novelo-Casanova, C. Chico and D. Ruíz (2016). Active Crustal Deformation in the Trans Mexican Volcanic Belt as Evidenced by Earthquakes During the last 450 Years, *J. Geophys. Res.*, Submitted.
- Sunye-Puchol, I., Lacan, P., Ortuño, M., Villamor, P., Audin, L., Zúñiga, F. R. & Lawton, T. F. (2015). La falla San Mateo: nuevas evidencias paleosismológicas de fallamiento activo en el graben de Acambay, México, *Revista Mexicana de Ciencias Geológicas*, 32(3), 361-376.
- Suter, M., Quintero, O., and Johnson, C. A. (1992). Active faults and state of stress in the central part of the Trans-Mexican Volcanic Belt, Mexico 1. The Venta de Bravo Fault. *Journal of Geophysical Research: Solid Earth* (1978–2012), 97(B8), 11983-11993.
- Suter, M., Quintero-Legorreta, O., López-Martínez, M., Aguirre-Díaz, G., and Farrar, E. (1995). The Acambay graben: active intraarc extension in the Trans-Mexican volcanic belt, Mexico. *Tectonics*, 14(6), 1245-1262.
- Suter, M., Carrillo-Martínez, M., and Quintero-Legorreta, O. (1996). Macro seismic study of shallow earthquakes in the central and eastern parts of the trans-Mexican volcanic belt, Mexico. *Bulletin of the Seismological Society of America*, 86(6), 1952-1963.
- Suter, M., Martínez, M. L., Legorreta, O. Q., and Martínez, M. C. (2001). Quaternary intra-arc extension in the central Trans-Mexican volcanic belt. *Geological Society of America Bulletin*, 113(6), 693-703.
- Suter M., 2015. The AD 1567 Mw 7.2 Ameca, Jalisco, Earthquake (Western Trans-Mexican Volcanic Belt): Surface Rupture Parameters, Seismogeological Effects, and Macro seismic Intensities from Historical Sources. *Bulletin of the Seismological Society of America*, 105(2A), pp.646-656.
- Urbina, F., and H. Camacho (1913). La zona megasísmica Acambay-Tixmadejé, Estado de México, conmovida el 19 de noviembre de 1912, Instituto Geológico de México, Boletín, 32, 125p.
- Ziagos, J. P., Blackwell, D. D., and Mooser, F. (1985). Heat flow in southern Mexico and the thermal effects of subduction. *Journal of Geophysical Research: Solid Earth*, 90(B7), 5410-5420.
- Zúñiga, F.R. y M.Wyss, (1995). Inadvertent changes in magnitude reported in earthquake catalogs: Influence on b-value estimates, *Bull. Seismol. Soc. Am.*, V.85, 1858-1866.
- Zúñiga, F. R., Pacheco, J. F., Guzmán-Speziale, M., Aguirre-Díaz, G. J., Espindola, V. H., and Nava, E. (2003). The Sanfandila earthquake sequence of 1998, Queretaro, Mexico: activation of an undocumented fault in the northern edge of central Trans-Mexican Volcanic Belt. *Tectonophysics*, 361(3), 229-238.
- Zúñiga, F.R., M. Reyners and P. Villamor, (2005). Temporal variations of the earthquake data in the catalogue of seismicity of New Zealand, *Bull. New Zeal. Soc. Earthquake Eng.*, 38, 87-107.
- Zúñiga, F. R., Ortuño M., A. Figueroa-Soto, (2012). Incorporation Of Paleoseismological Data in the Calculation of the Seismic Hazard: An Example in Central Mexico. *Proceedings of the 3rd INQUA-IGCP 567 International Workshop on Earthquake Geology, Palaeoseismology and Archaeoseismology, Morelia México*.
- Zúñiga, F.R., G. Suárez, A. Figueroa and A. Mendoza (2016). A first order seismotectonic regionalization of Mexico for seismic hazard and risk estimation, *J. Seismol.*, in review.